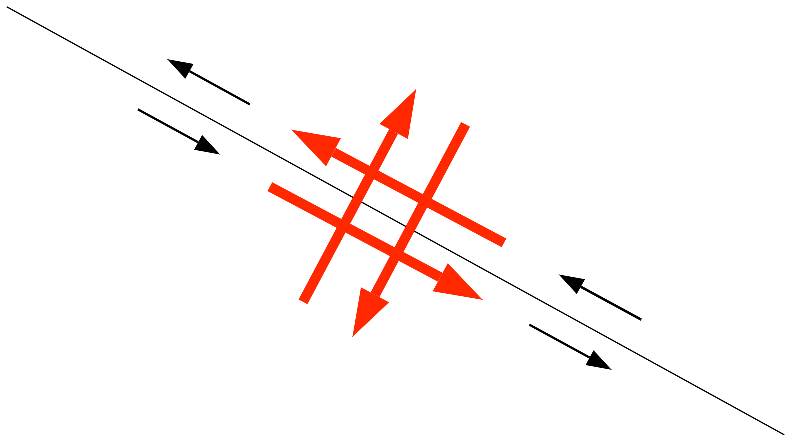
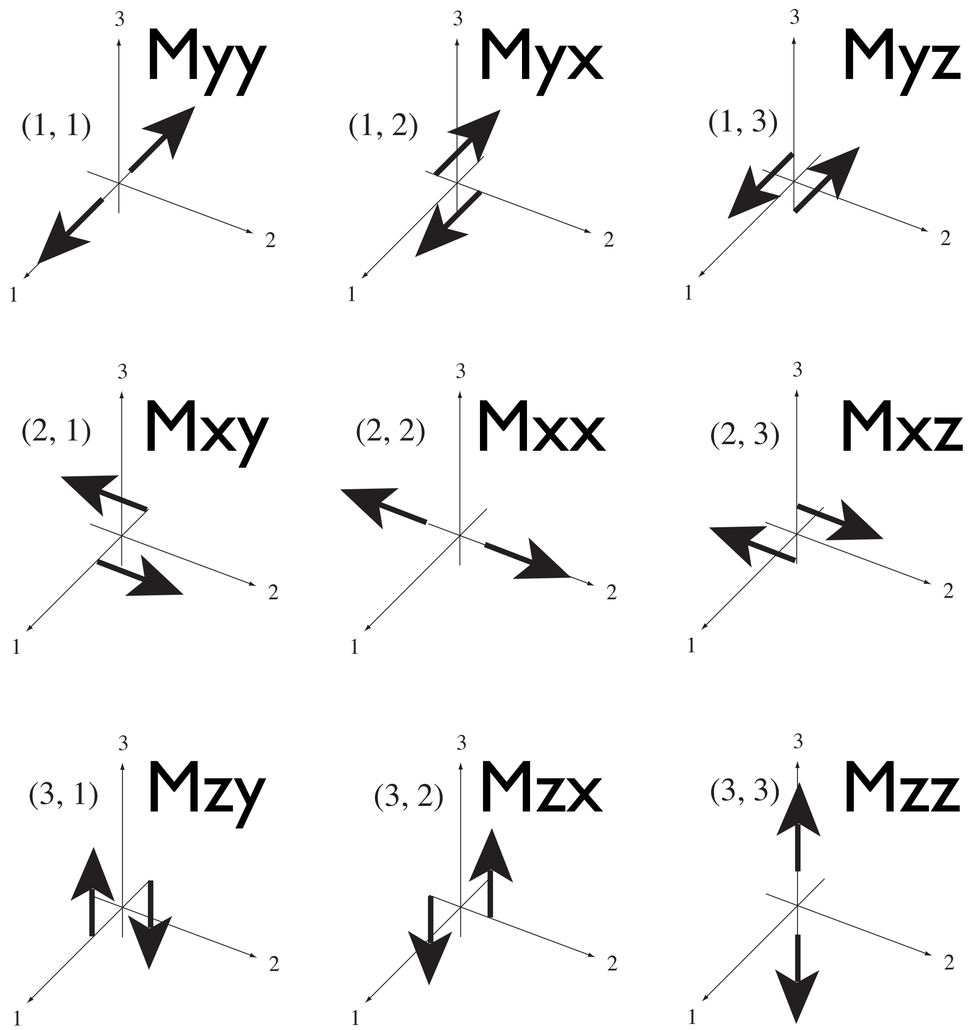
1. Moment-tensor analysis using global data
2. The Global CMT catalog
3. Using calibration information in waveform analysis
4. Data quality control using signals
5. Data quality control using noise
6. Finding interesting things in the noise

Faulting force model



The elastic stress release in an earthquake is described by a double couple of forces

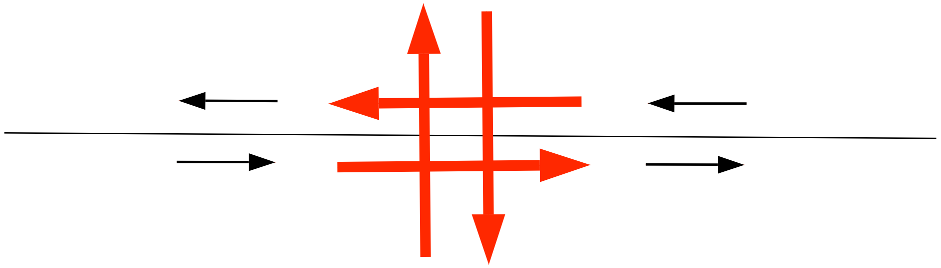
The nine dipoles of the seismic moment tensor



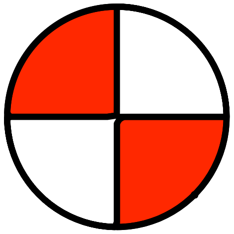
(Aki and Richards, 2002)

But, Mxy=Myx, Myz=Mzy, Mxz=Mzx

Mxy =



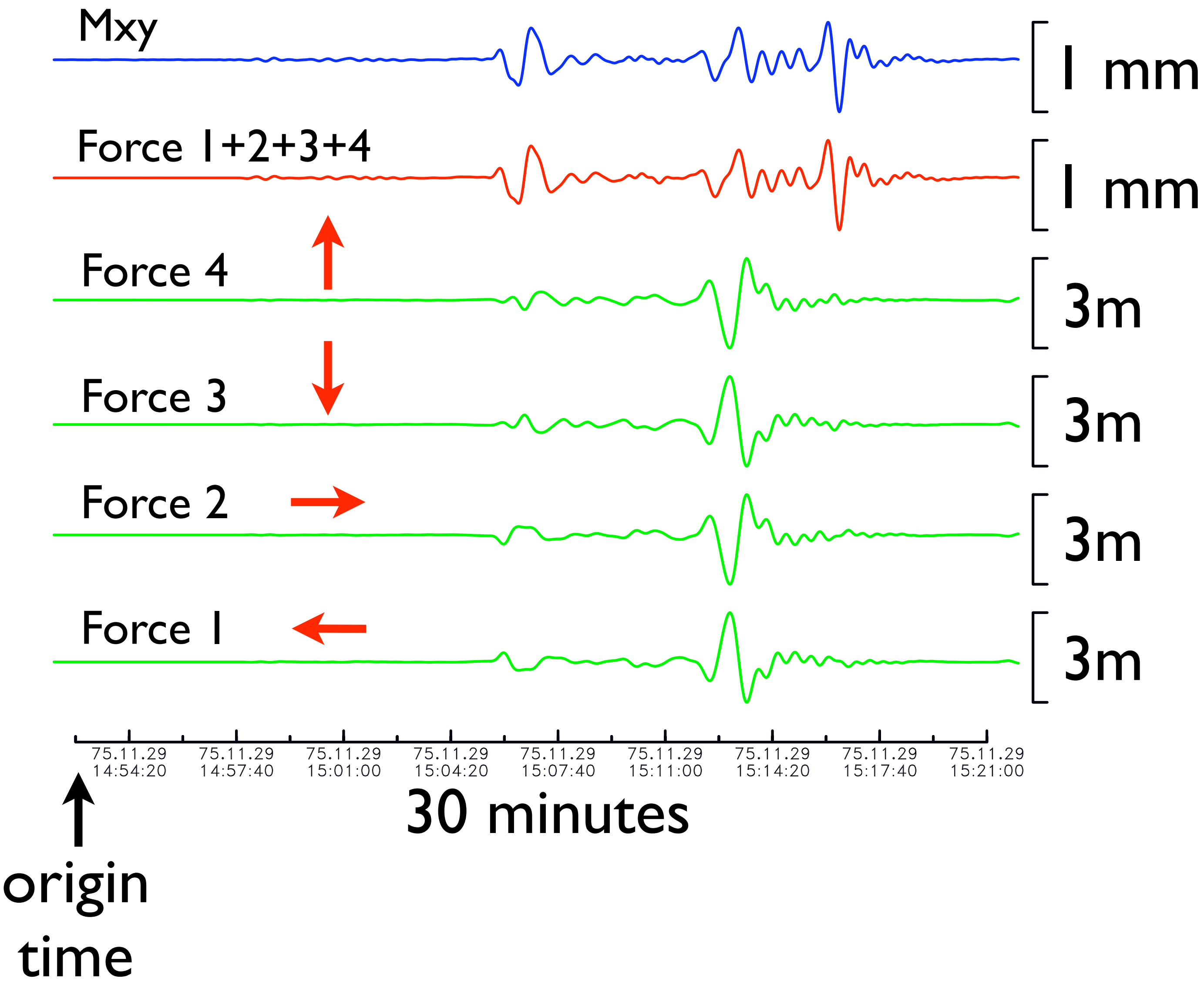
=



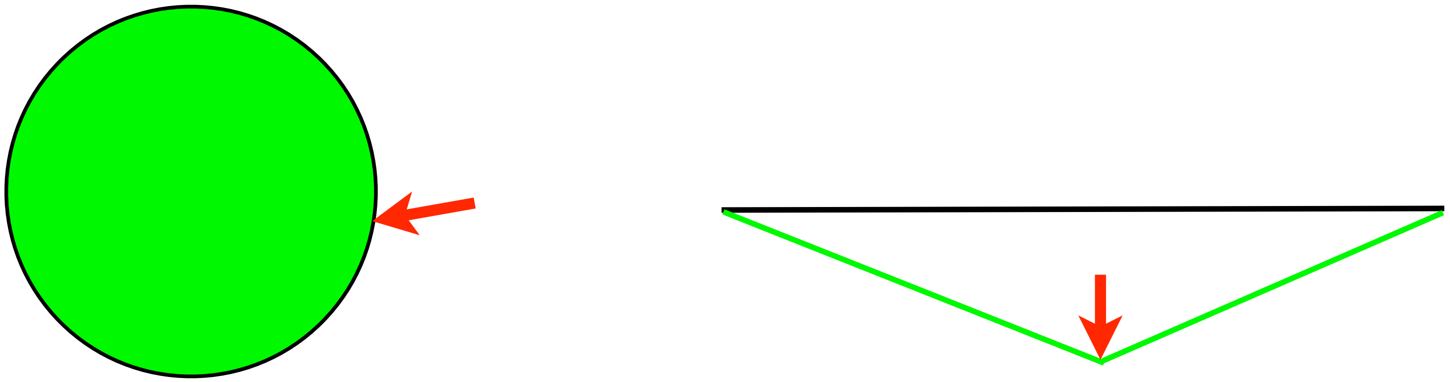
for example,

10^{28} dyne-cm = 10^{24} dyne x 10000 cm

Calculated force seismograms (6000 km distance)



The vibrations caused by a force acting on or in the Earth can be modeled by summation of Earth's normal modes



$$u(\boldsymbol{x},t) = \sum_k \left[1 - \exp\left[-\alpha_k(t-t_s)\right]\cos\omega_k(t-t_s)\right] \boldsymbol{f}\cdot\boldsymbol{w}^{(k)}(\boldsymbol{x}_s)\boldsymbol{s}_k(\boldsymbol{x})$$

where $\boldsymbol{f}$ is the force vector and $\boldsymbol{w}^k$ is the displacement of the k-*th* mode.

Moment-tensor analysis by waveform fitting

(Observed seismogram)/(Instrument response) x Filter = Observed waveform

(Synthetic displacement seismogram) x Filter = Model waveform

Model waveform depends on:
1. Earthquake parameters
2. Earth structure

If the Earth structure and the earthquake location are known, the

Model waveform depends only on the six elements of the moment tensor,

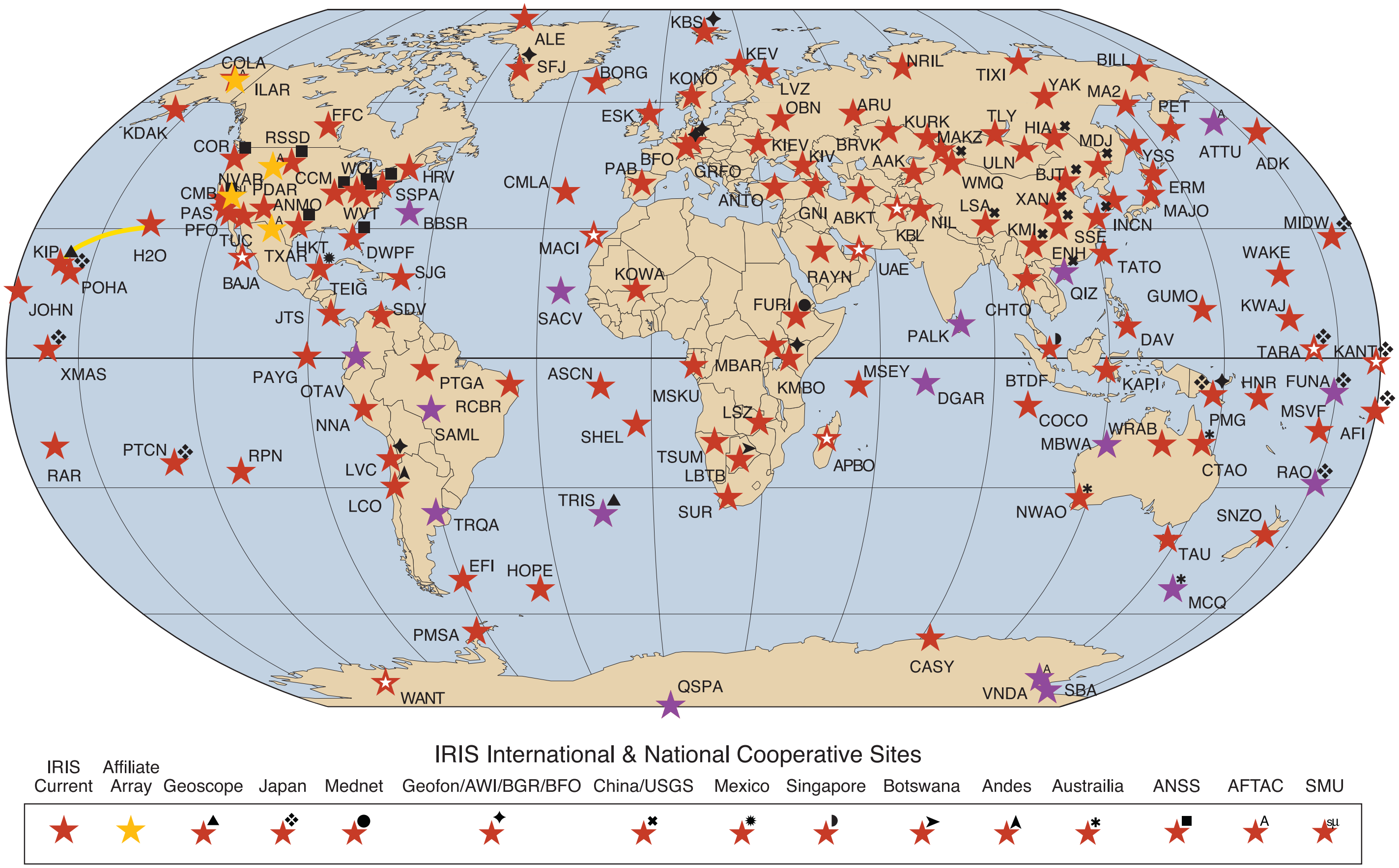
M_{xx}, M_{yy}, M_{zz}, M_{xy}, M_{xz}, and M_{yz}

Minimize the difference [Observed waveform - Model waveform]^2

with respect to the moment tensor elements.

Detection and analysis of large earthquakes:

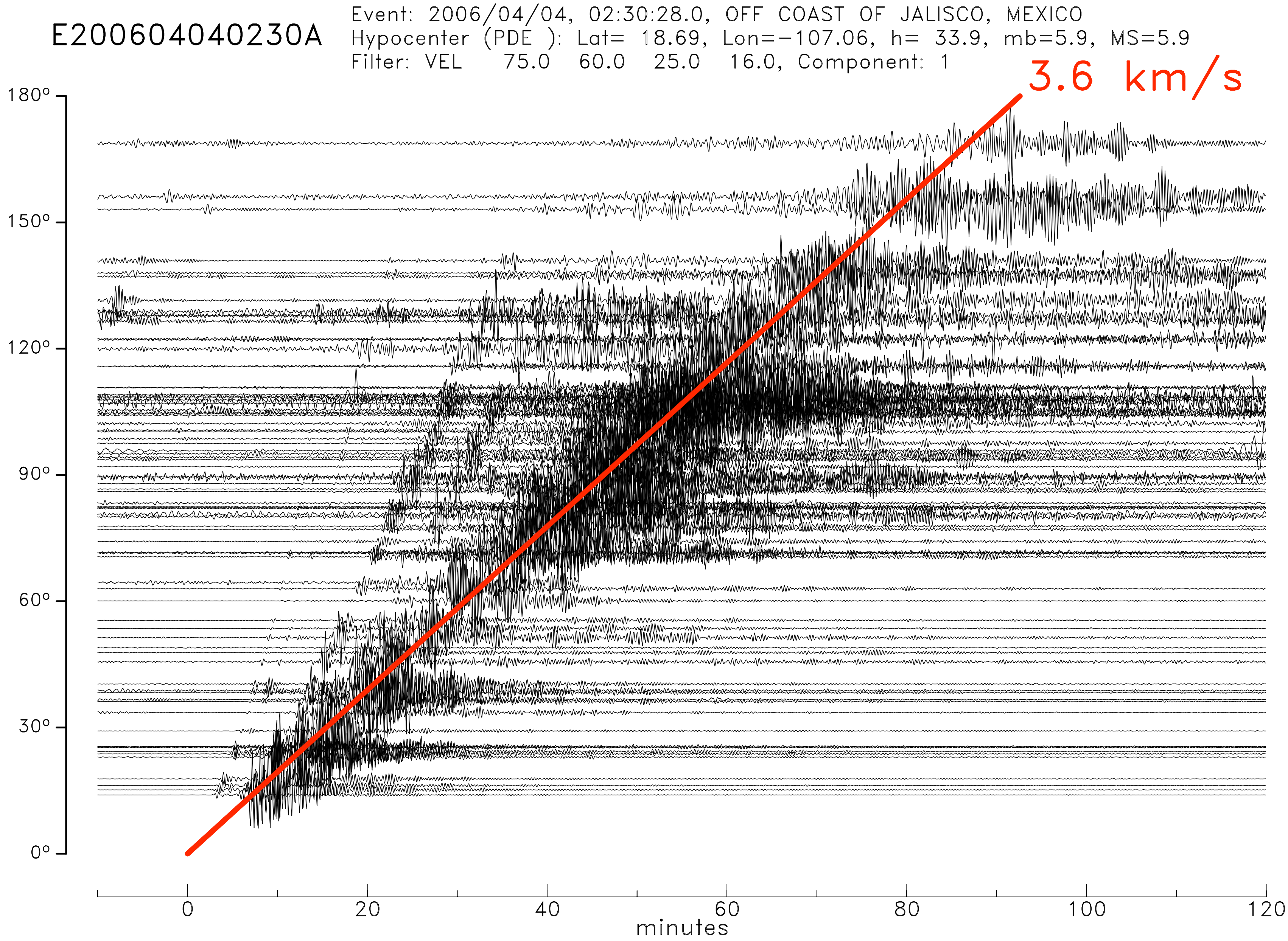
GLOBAL SEISMOGRAPHIC NETWORK



Current GSN station coverage of Earth is shown as of August 2005. Sites added in the past five years are noted in purple (stations) and orange (arrays). Sites planned to be completed are noted with white stars. Cooperative sites are indicated by symbols on the upper right "shoulder" of the stars.

STS-1 Seismometer
at Harvard, Mass.

Global network record section for an earthquake off the coast of Jalisco, Mexico



Moment-tensor analysis by waveform fitting

(Observed seismogram)/(Instrument response) x Filter = Observed waveform

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If the Earth structure and the earthquake location are known, the

Model waveform depends only on the six elements of the moment tensor,

M_{xx}, M_{yy}, M_{zz}, M_{xy}, M_{xz}, and M_{yz}

Minimize the difference $[\text{Observed waveform - Model waveform}]^2$

with respect to the moment tensor elements.

Seismogram Modeling

The k-*th* seismogram in a data set for a given earthquake can be represented by:

$$u_k(\boldsymbol{r}, t) = \sum_{i=1}^{N} \psi_{ik}(\boldsymbol{r}_0, \boldsymbol{r}, t) f_i$$

where ψ_{ik} are the excitation kernels and f_i are independent parameters of the source model.

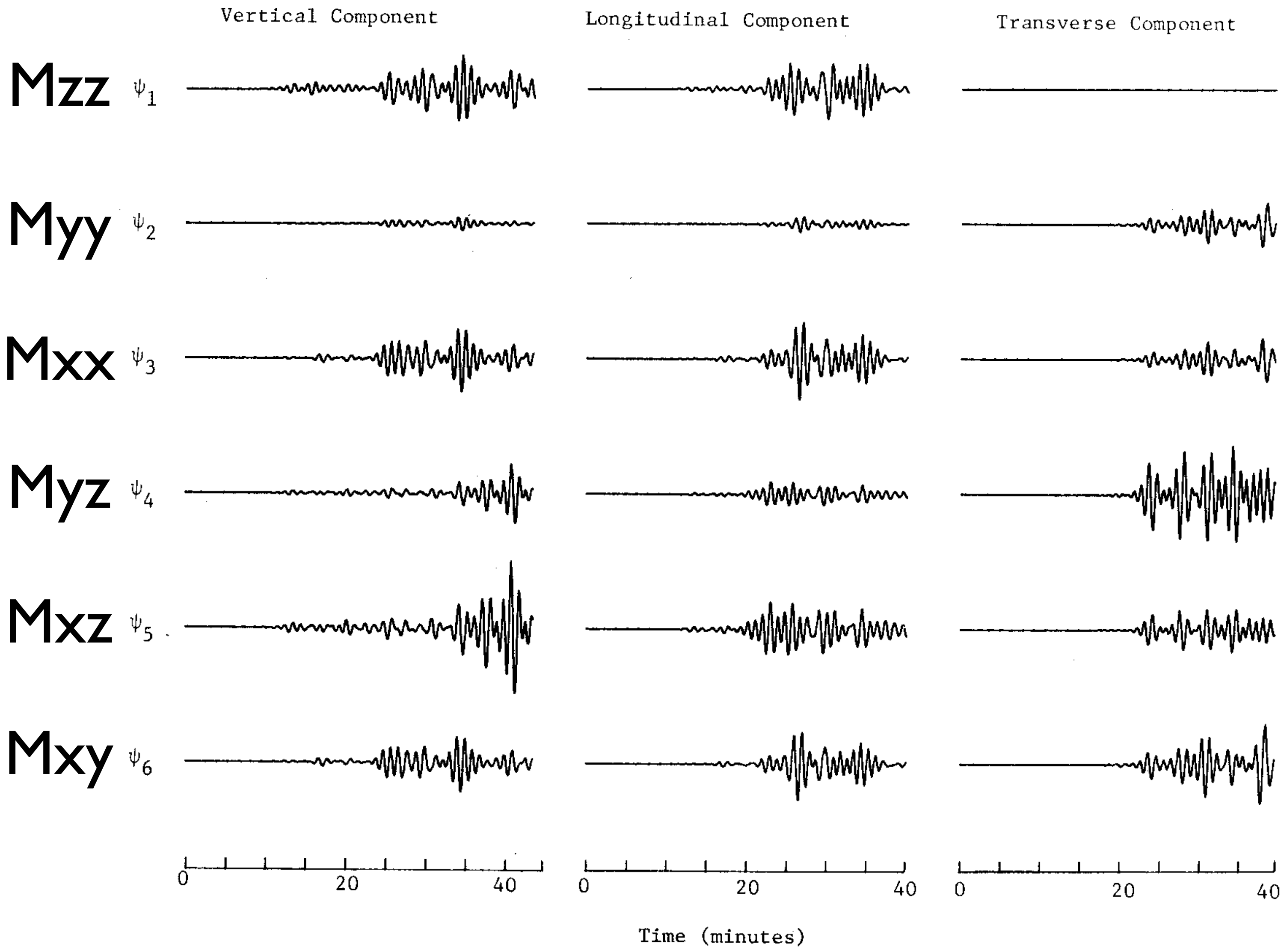
f_1 = Mzz, f_2 =Myy, etc.; N=6

Seismogram Synthesis for a Moment-Tensor Source

The seismic displacement field can be calculated by superposition of the normal modes of the Earth (Gilbert, 1971):

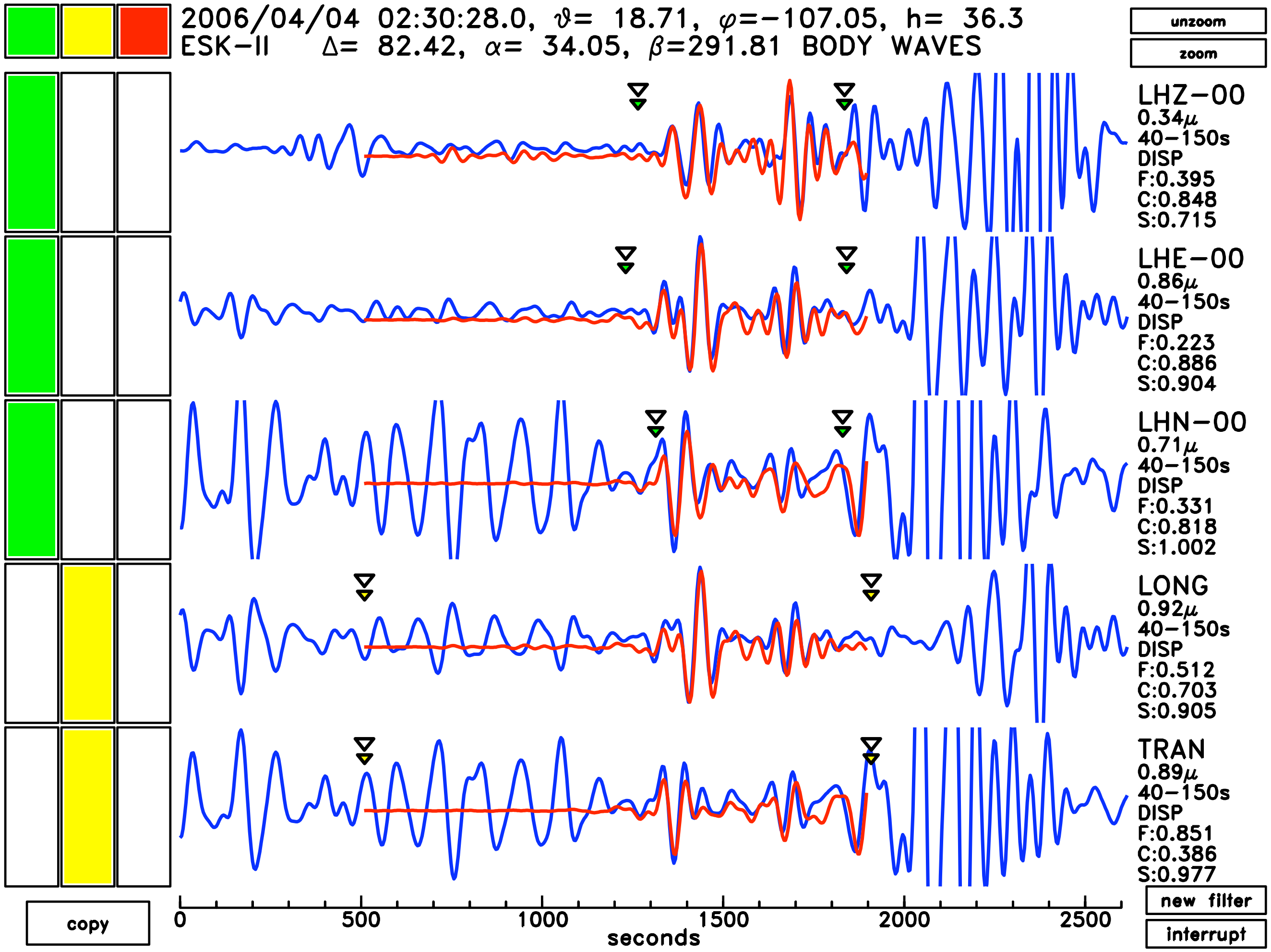
$$u(\boldsymbol{x},t) = \sum_k \left[1 - \exp\left[-\alpha_k(t-t_s)\right]\cos\omega_k(t-t_s)\right] \boldsymbol{M} : \boldsymbol{e}^{(k)}(\boldsymbol{x}_s)\boldsymbol{s}_k(\boldsymbol{x})$$

where α_k is the decay constant of and $\boldsymbol{e}^k$ is the strain tensor in the k-*th* mode; $\boldsymbol{s}_k$ is the eigenfunction of the k-*th* mode; and $\boldsymbol{M}$ is the seismic moment tensor.



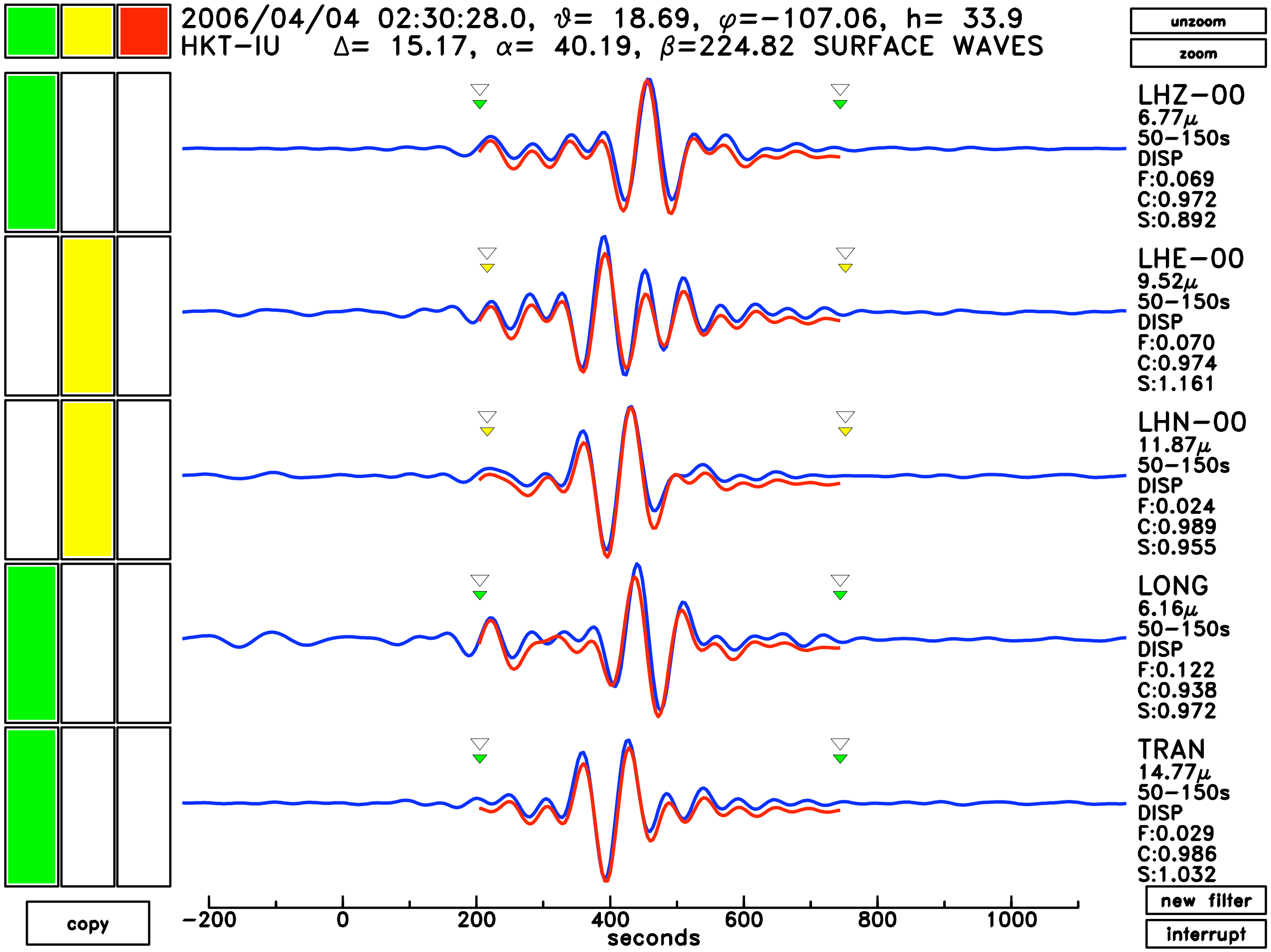
Excitation kernels for deep earthquake (580 km)

Fit to seismograms:
Body waves at Eskdalemuir, Scotland



blue - data ; red - model

Fit to seismograms:
Surface waves at Hockley, Texas



```
2006/04/04 02:30:28.0, ϑ= 18.69, φ=-107.06, h= 33.9
HKT-IU    Δ= 15.17, α= 40.19, β=224.82 SURFACE WAVES
```

```
LHZ-00
6.77μ
50-150s
DISP
F:0.069
C:0.972
S:0.892

LHE-00
9.52μ
50-150s
DISP
F:0.070
C:0.974
S:1.161

LHN-00
11.87μ
50-150s
DISP
F:0.024
C:0.989
S:0.955

LONG
6.16μ
50-150s
DISP
F:0.122
C:0.938
S:0.972

TRAN
14.77μ
50-150s
DISP
F:0.029
C:0.986
S:1.032
```

blue - data ; red - model

Estimation of the Source Parameters

For a point source, the elements f_i can be estimated by solving $\mathbf{A} \cdot \mathbf{f} = \mathbf{b}$, where:

$$A_{ij} = \sum_k \int_{t_{k_1}}^{t_{k_2}} \psi_{ik} \psi_{jk} dt \ ; \ b_j = \sum_k \int_{t_{k_1}}^{t_{k_2}} u_k \psi_{jk} dt.$$

This procedure requires that the position of the source $(\boldsymbol{r}_0, t_0)$ be known.

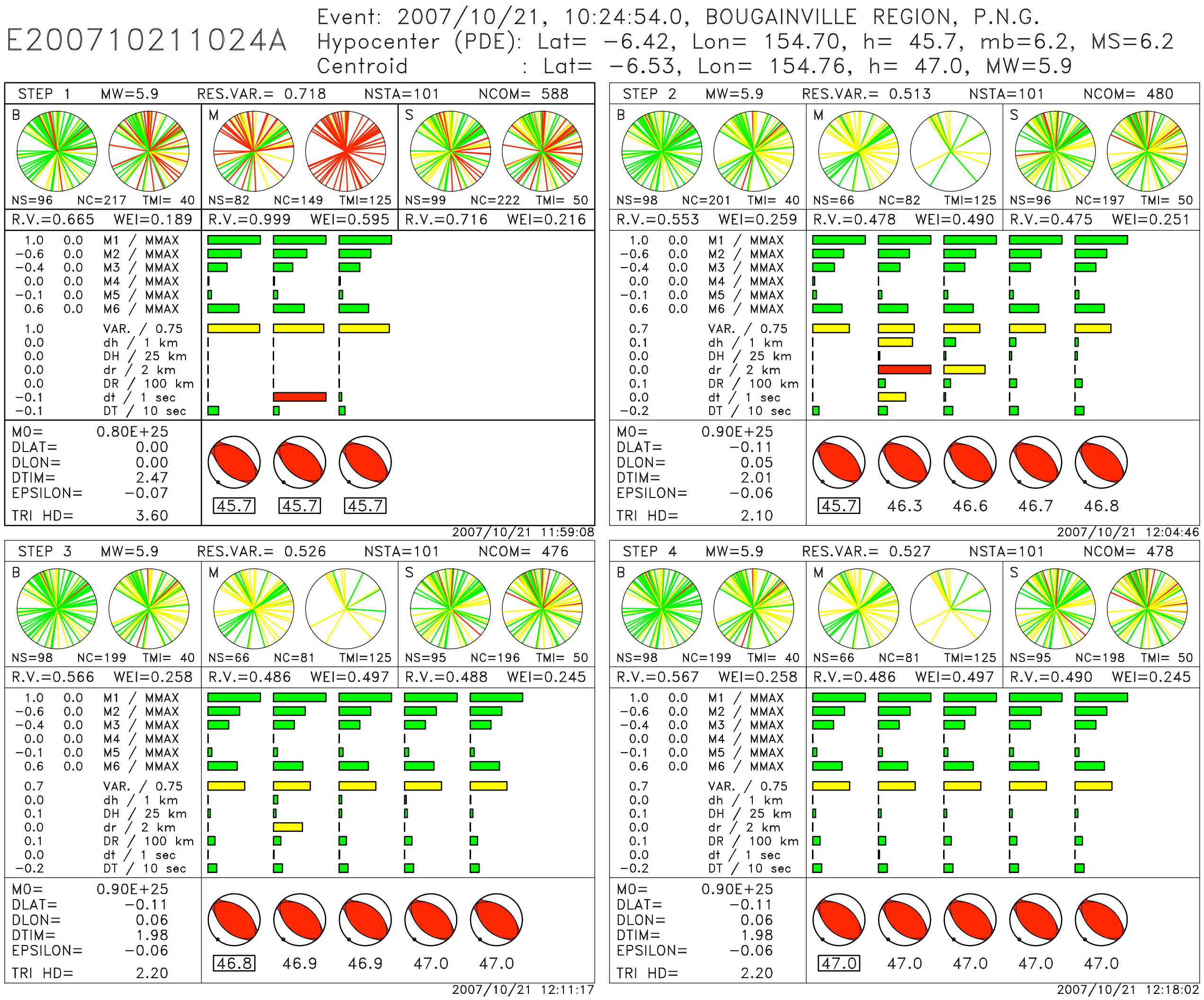
Solution for the Source Centroid

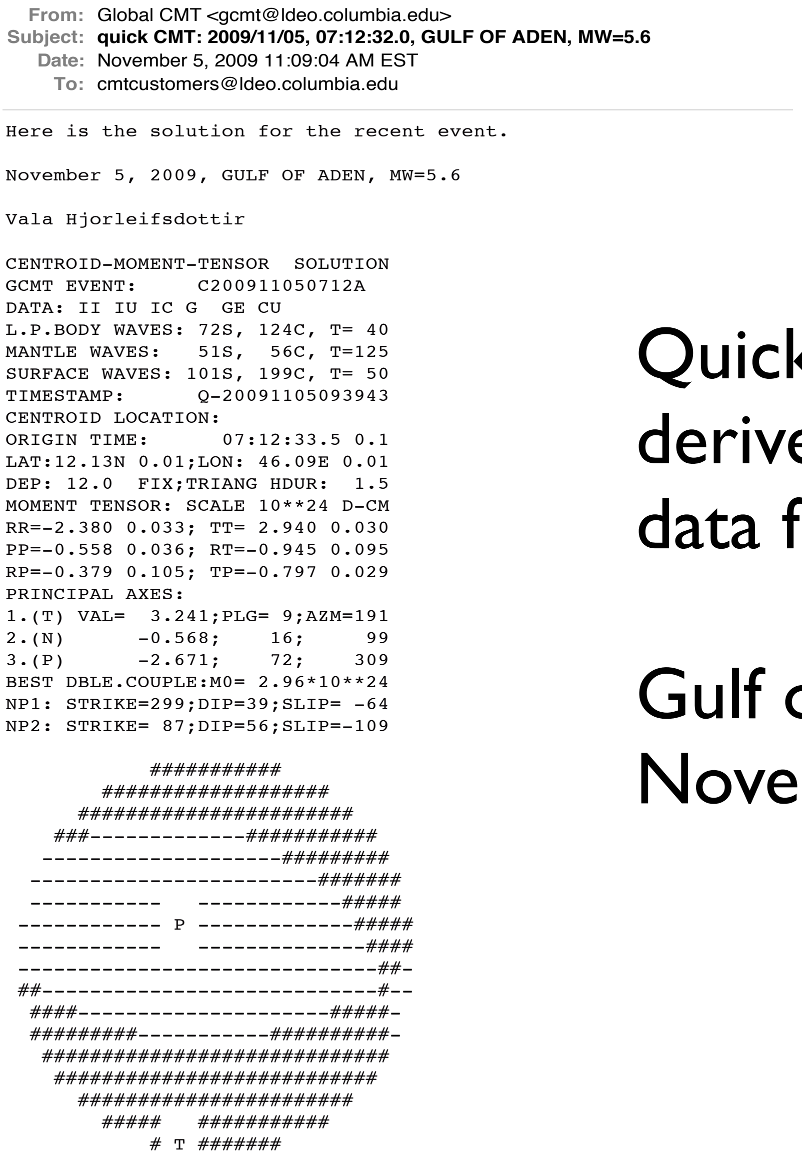
The earthquake centroid can be determined simultaneously with the source model parameters by expansion of the equations of condition to allow for a perturbation in the location of the source (Dziewonski, Chou and Woodhouse, 1981):

$$u_k = u_k^{(0)} + \{\psi_{ki,j}^{(0)} \cdot \delta x_j - \psi_{ki,t}^{(0)} \cdot \delta t_0\} \cdot f_i^{(0)} + \psi_{ki}^{(0)} \cdot \delta f_i \; ;$$

where the superscript (0) indicates parameters determined for the starting location. The problem can then be solved iteratively.

Iterative procedure for moment-tensor source converges nicely





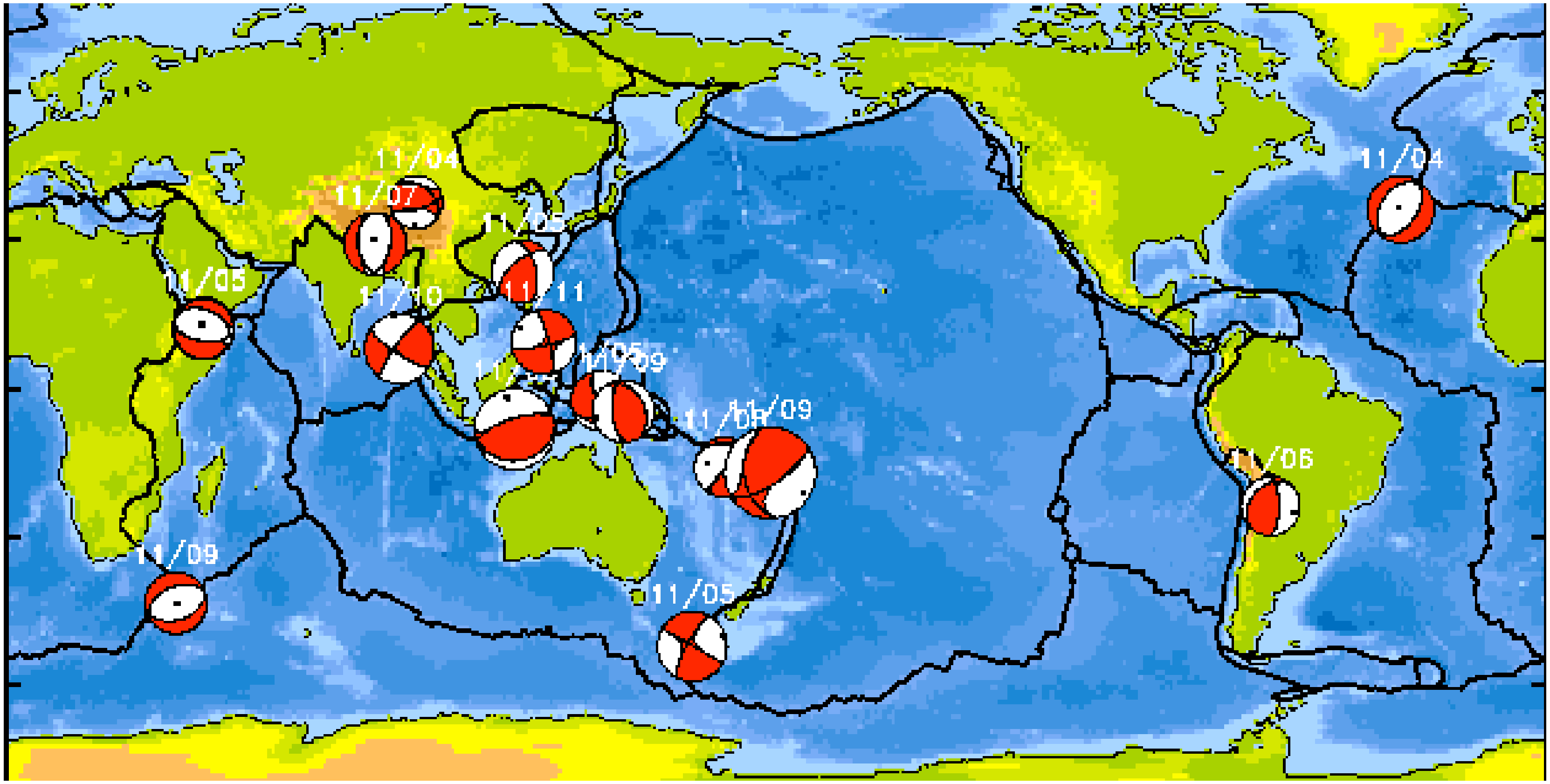
From: Global CMT <gcmt@ldeo.columbia.edu>
Subject: **quick CMT: 2009/11/05, 07:12:32.0, GULF OF ADEN, MW=5.6**
Date: November 5, 2009 11:09:04 AM EST
To: cmtcustomers@ldeo.columbia.edu

```
Here is the solution for the recent event.

November 5, 2009, GULF OF ADEN, MW=5.6

Vala Hjorleifsdottir

CENTROID-MOMENT-TENSOR  SOLUTION
GCMT EVENT:     C200911050712A
DATA: II IU IC G  GE CU
L.P.BODY WAVES: 72S, 124C, T= 40
MANTLE WAVES:   51S,  56C, T=125
SURFACE WAVES: 101S, 199C, T= 50
TIMESTAMP:      Q-20091105093943
CENTROID LOCATION:
ORIGIN TIME:      07:12:33.5 0.1
LAT:12.13N 0.01;LON: 46.09E 0.01
DEP: 12.0  FIX;TRIANG HDUR:  1.5
MOMENT TENSOR: SCALE 10**24 D-CM
RR=-2.380 0.033; TT= 2.940 0.030
PP=-0.558 0.036; RT=-0.945 0.095
RP=-0.379 0.105; TP=-0.797 0.029
PRINCIPAL AXES:
1.(T) VAL=  3.241;PLG= 9;AZM=191
2.(N)      -0.568;    16;     99
3.(P)      -2.671;    72;    309
BEST DBLE.COUPLE:M0= 2.96*10**24
NP1: STRIKE=299;DIP=39;SLIP= -64
NP2: STRIKE= 87;DIP=56;SLIP=-109

            ###########
        ###################
      #######################
    ###-------------###########
   --------------------#########
  ------------------------#######
  -----------   ------------#####
 ------------ P -------------#####
 ------------   --------------####
 ------------------------------##-
 ##----------------------------#--
  ####---------------------#####-
  #########-----------##########-
   #############################
    ###########################
      #######################
        #####   ###########
            # T #######
```

Quick CMT solutions derived from real-time data from the GSN:

Gulf of Aden,
November 5, 2009, M=5.6

Two weeks of quick CMTs, November 2009



2. The Global CMT catalog

3. Using calibration information in waveform analysis

4. Data quality control using signals

5. Data quality control using noise

6. Finding interesting things in the noise

The Global CMT Project

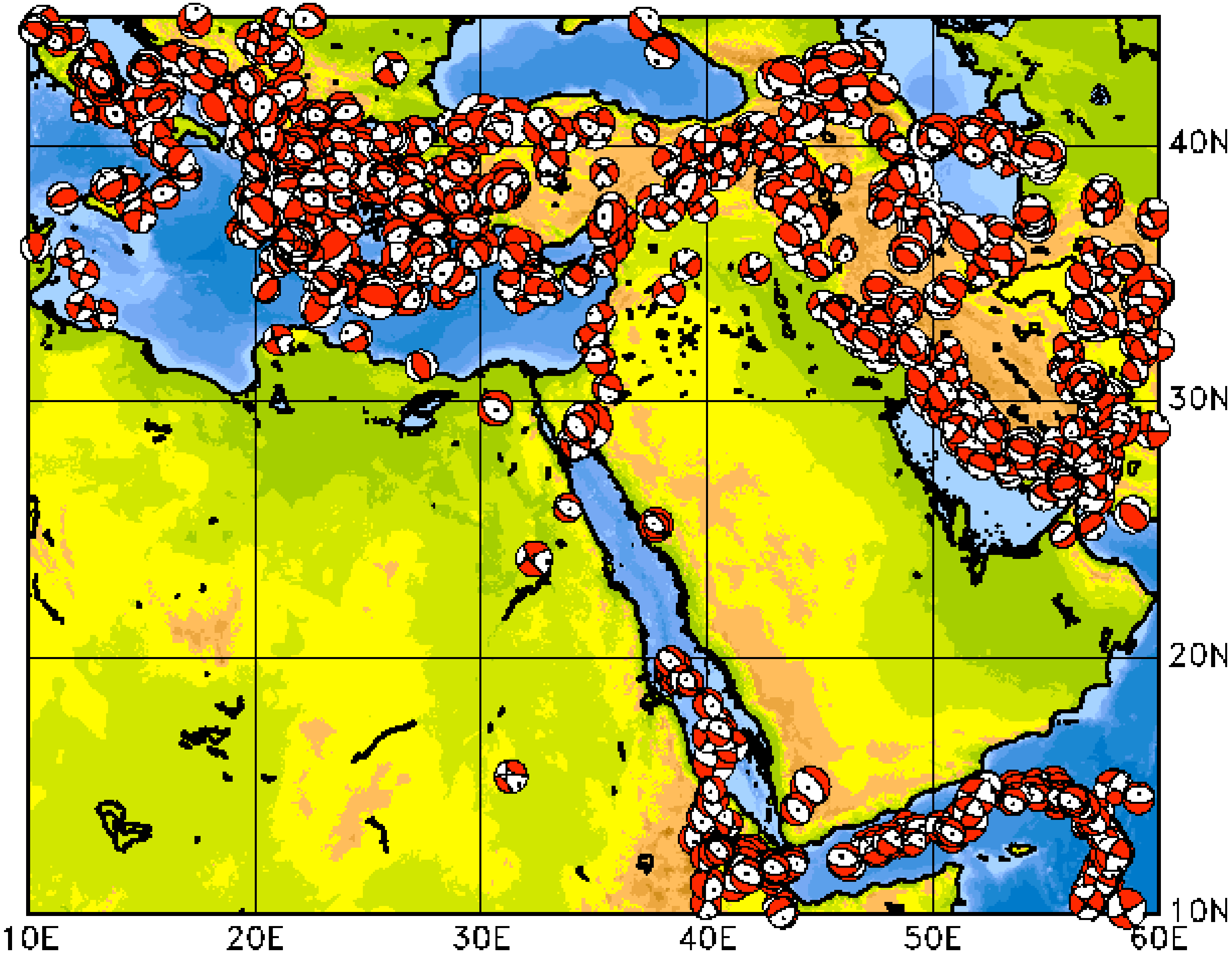
Project started in 1981 (A.M. Dziewonski et al.)

Goal is (now) to determine source parameters for
all earthquakes with M>5 worldwide

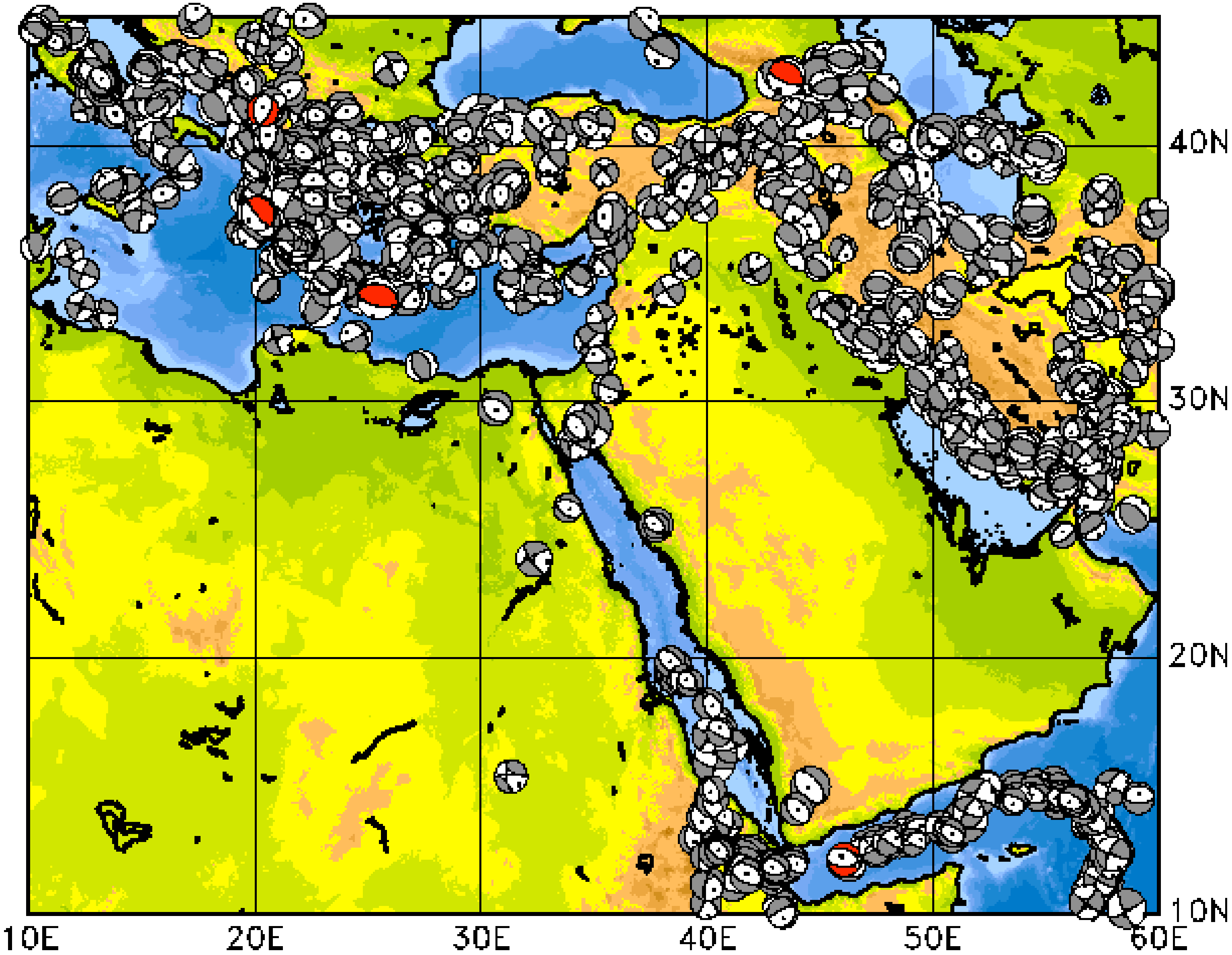
CMT catalog contains ~30,000 moment tensors
for the period 1976-2009

In 2006 the project moved from Harvard University
to Lamont-Doherty Earth Observatory
at Columbia University

Shallow earthquakes, 1976-2009

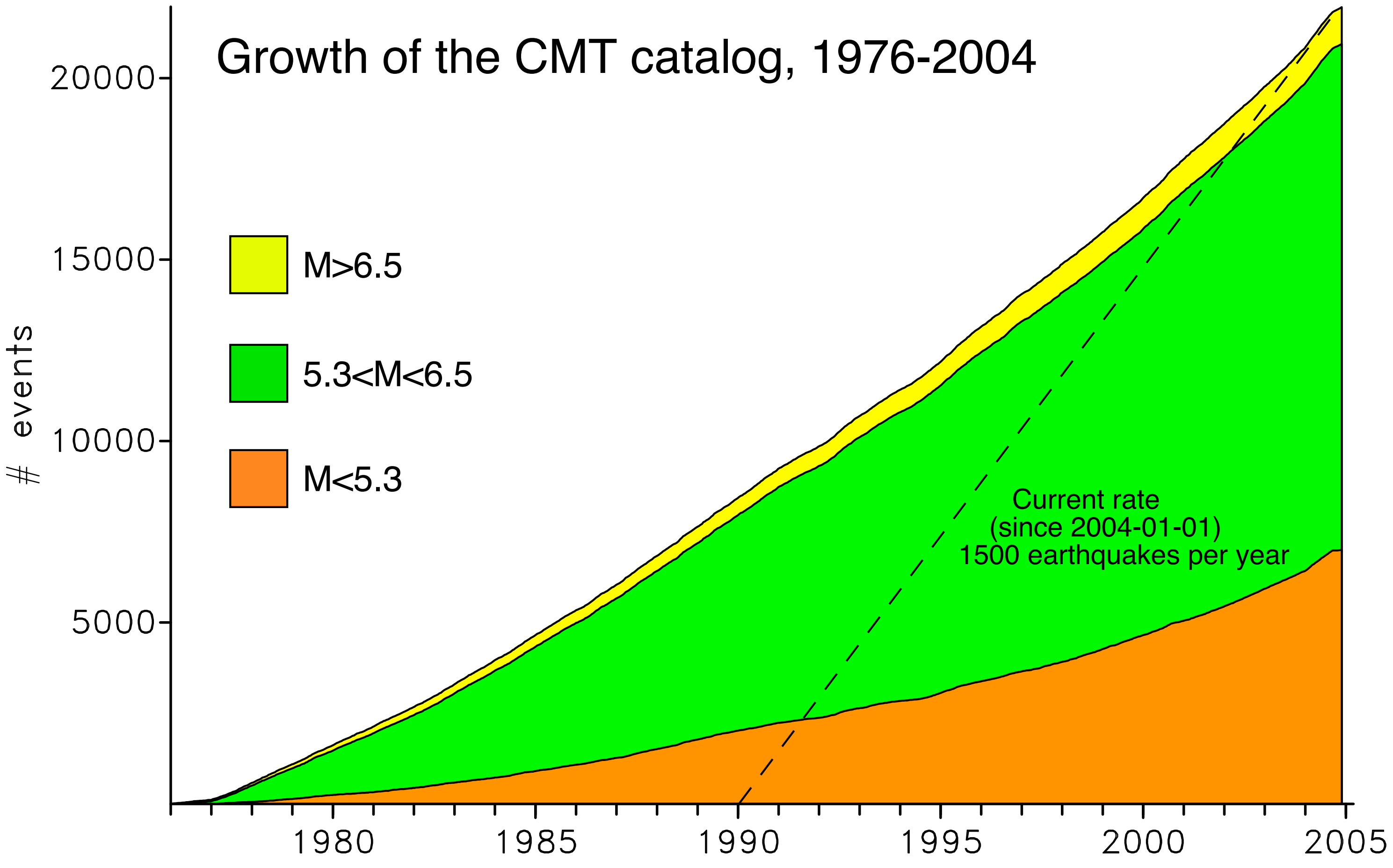


Shallow earthquakes, 1976-2009



red - earthquakes in last three months

We analyze 1500-2000 earthquakes per year



The CMT catalog can be accessed at
www.globalcmt.org

To receive Quick CMT solutions by email, send me an email at
ekstrom@ldeo.columbia.edu

3. Using calibration information in waveform analysis

4. Data quality control using signals

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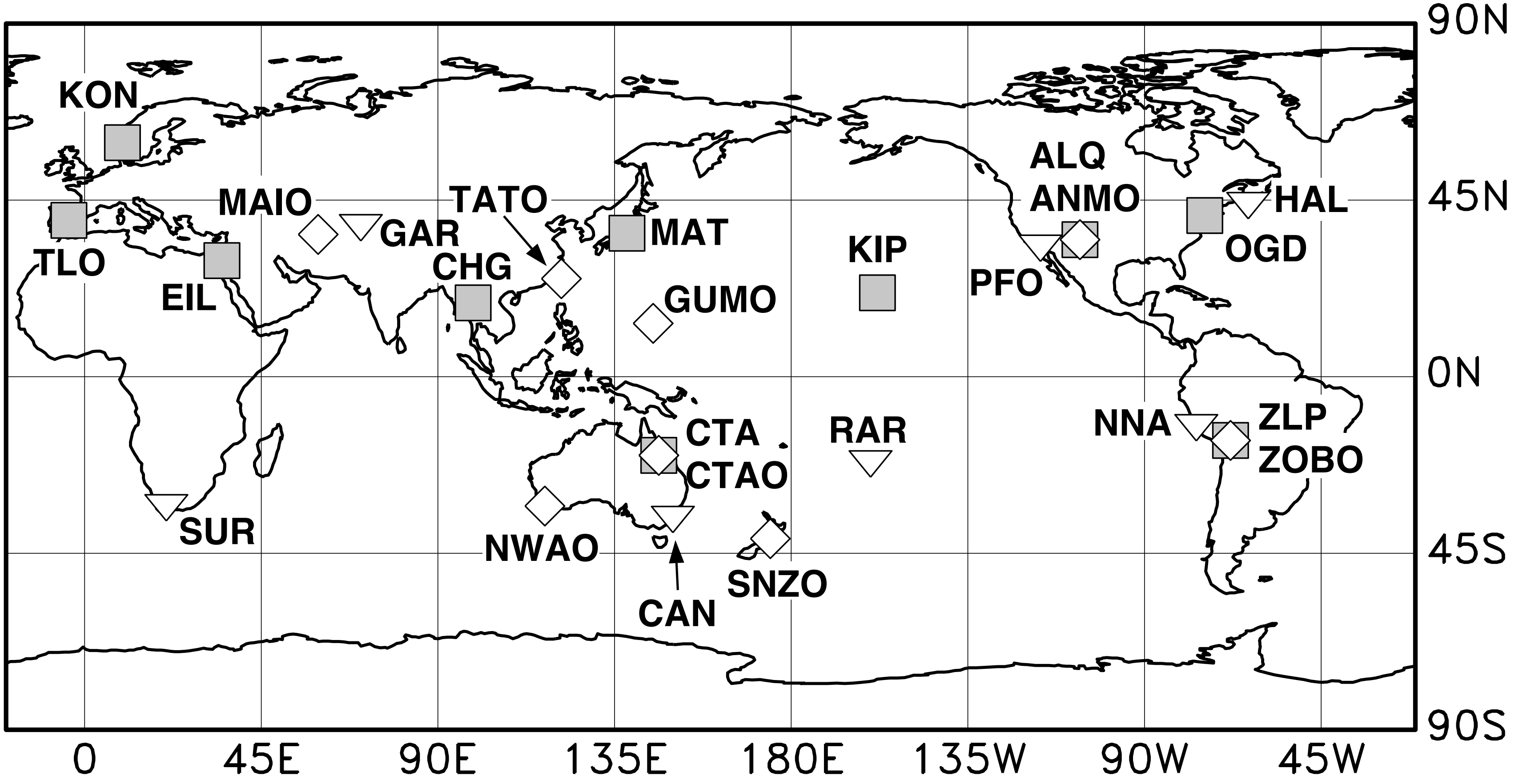
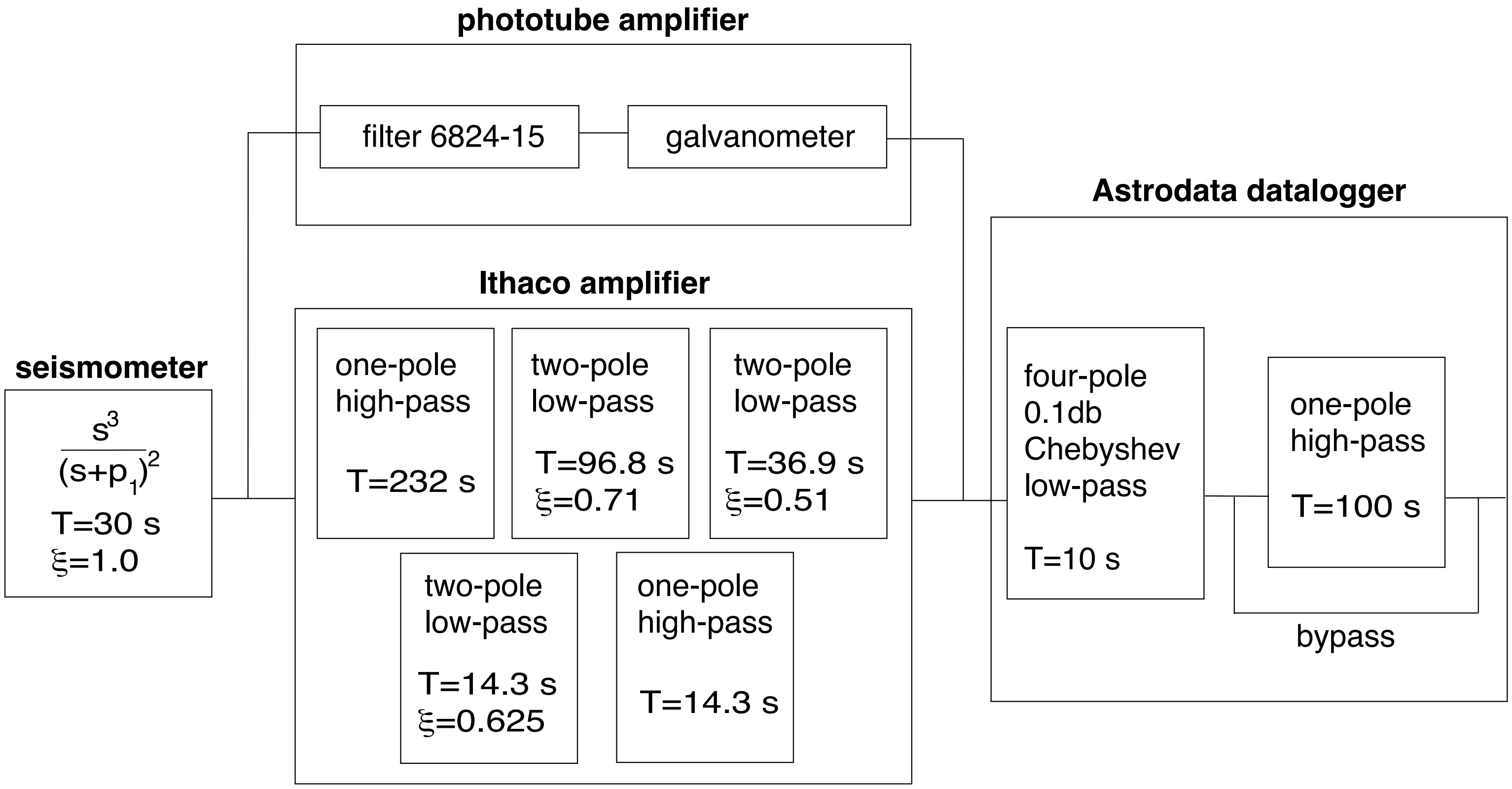
Quantitative waveform analysis requires
highly accurate instrument response information

The Global Digital Network in 1976



▨ High-Gain Long-Period (HGLP) network

HLGP seismometer and recording system



seismometer

$\frac{s^3}{(s+p_1)^2}$

T=30 s

ξ=1.0

phototube amplifier

filter 6824-15

galvanometer

Ithaco amplifier

one-pole high-pass

T=232 s

two-pole low-pass

T=96.8 s

ξ=0.71

two-pole low-pass

T=36.9 s

ξ=0.51

two-pole low-pass

T=14.3 s

ξ=0.625

one-pole high-pass

T=14.3 s

Astrodata datalogger

four-pole 0.1db Chebyshev low-pass

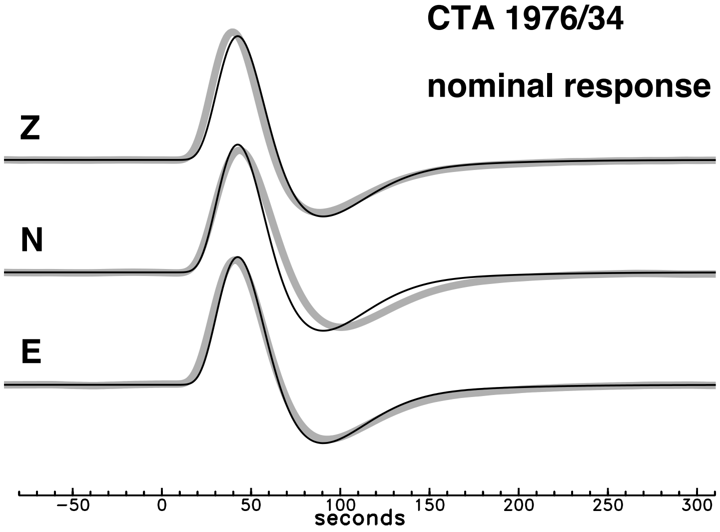
T=10 s

one-pole high-pass

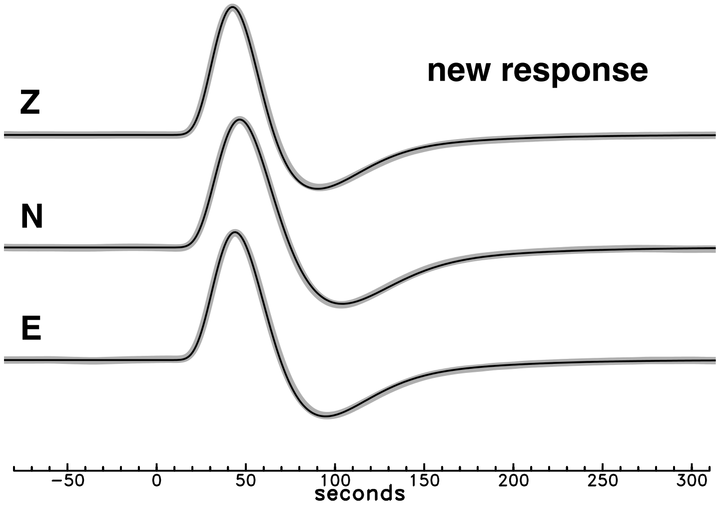
T=100 s

bypass

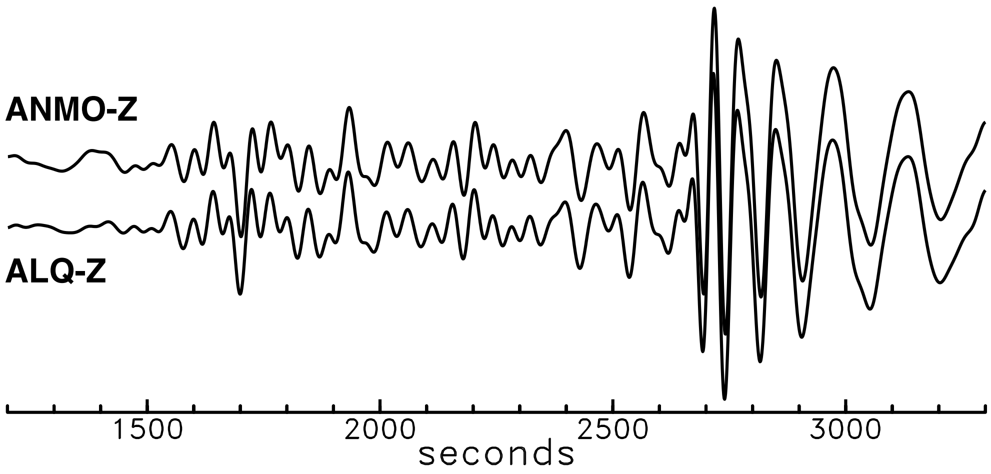
Original calibration pulses and pulses for nominal response



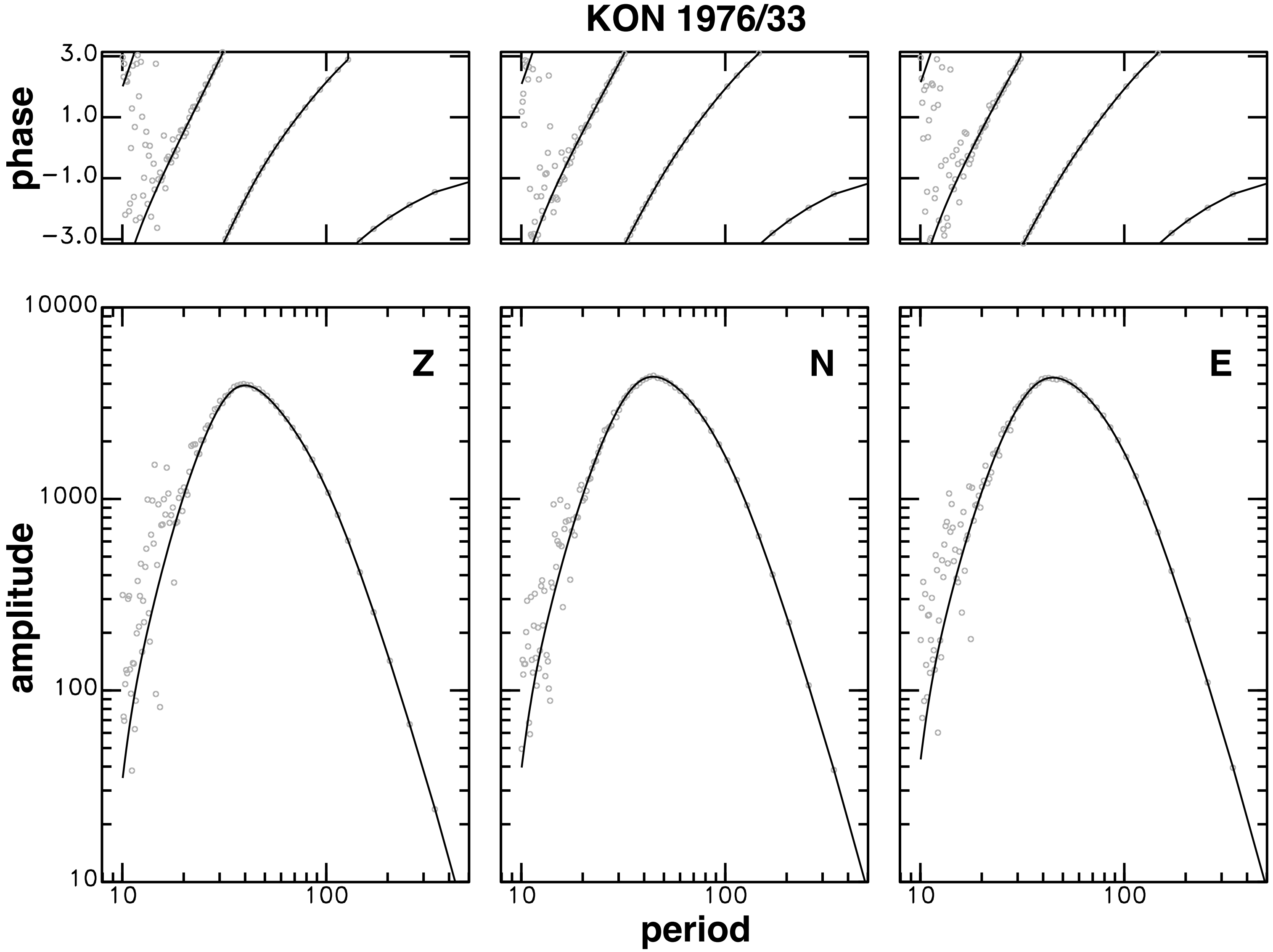
Original calibration pulses and pulses for new response after inversion



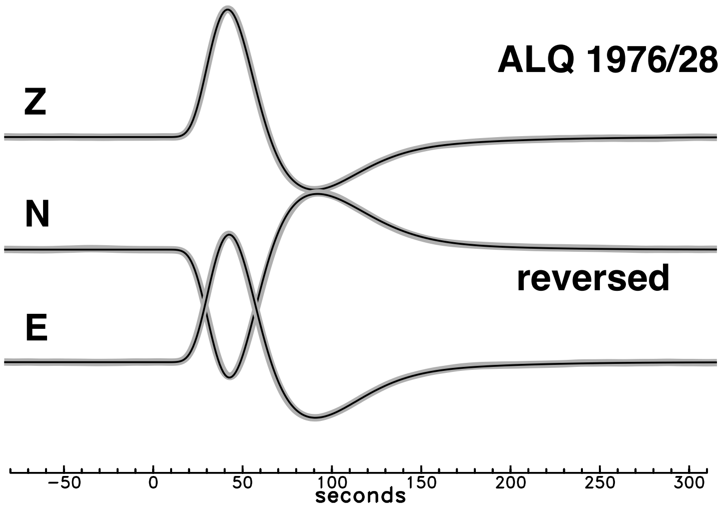
Comparison of waveforms after normalizing responses for two stations in the same location



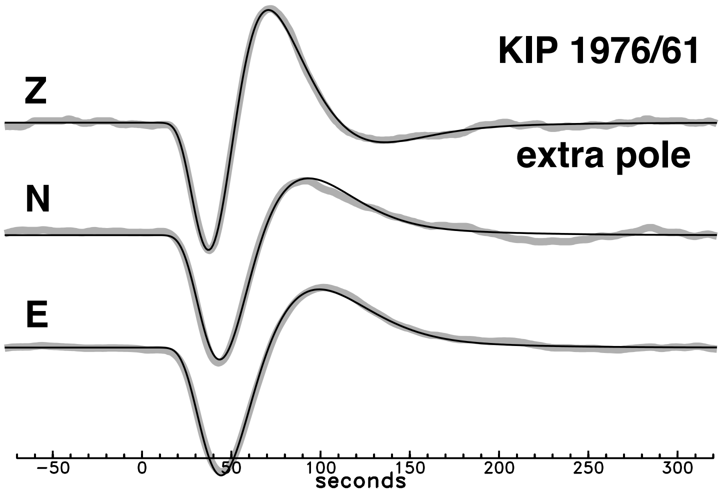
Check of new responses -- sine-wave calibrations



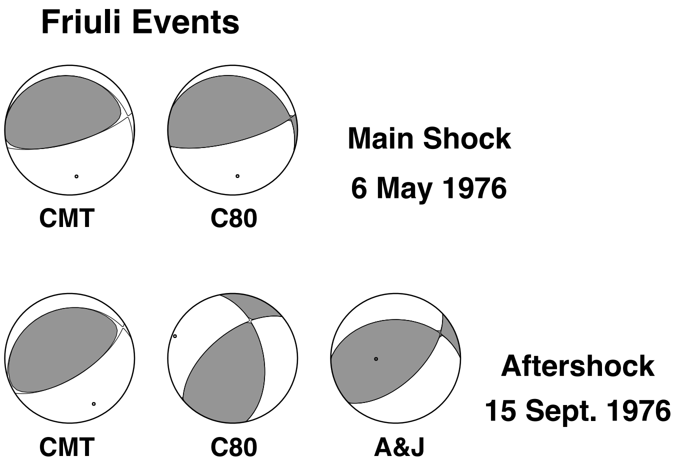
Some channels were reversed for some periods of time

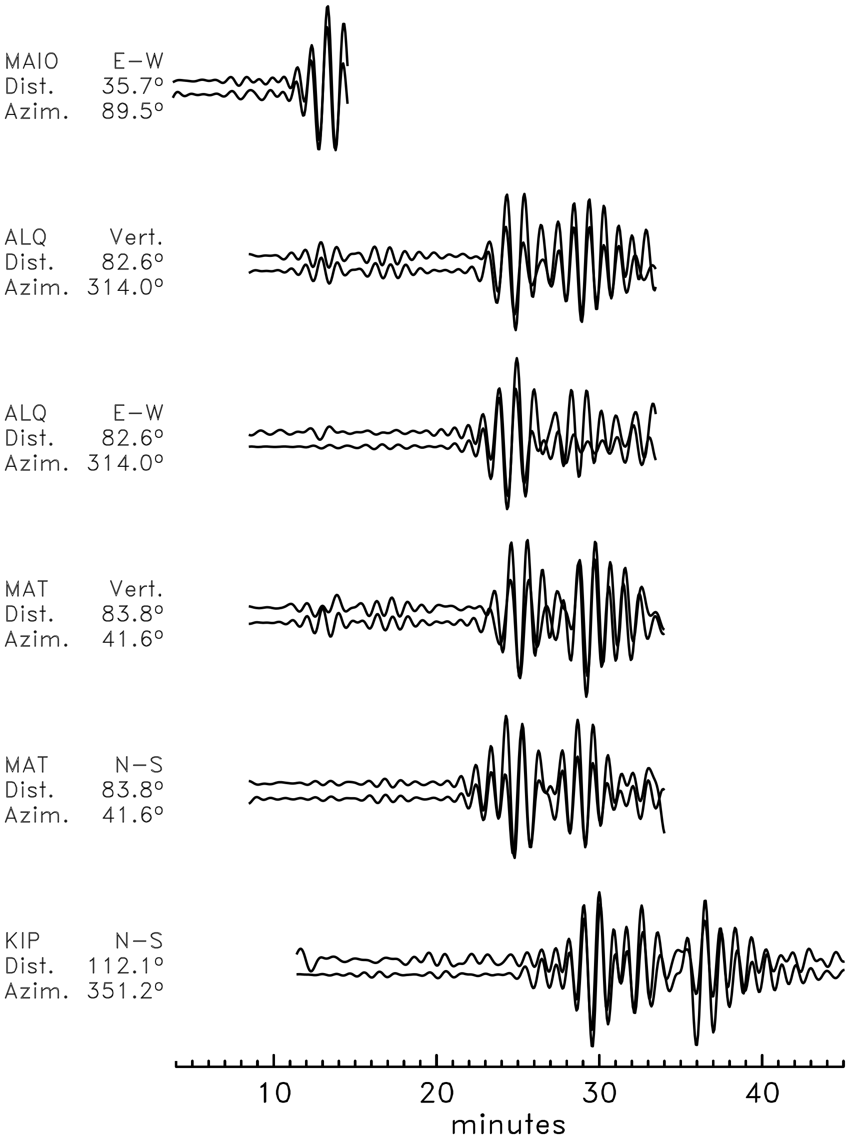


Some channels had extra filters for some periods of time



Waveform comparisons (observed and synthetic) after correcting seismograms using new responses: The 1976 Friuli earthquake





4. Data quality control using signals
5. Data quality control using noise
6. Finding interesting things in the noise

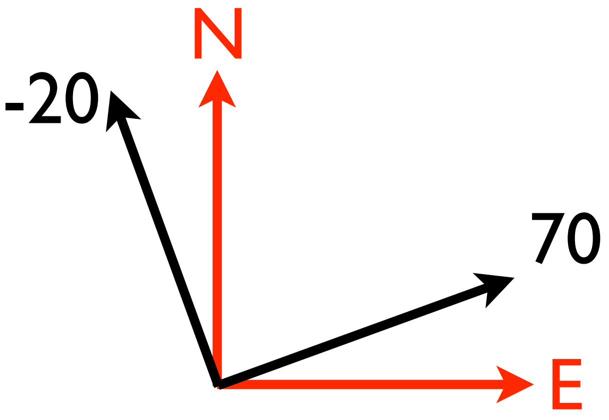
4a. Sensor orientation

4b. Sensor response stability

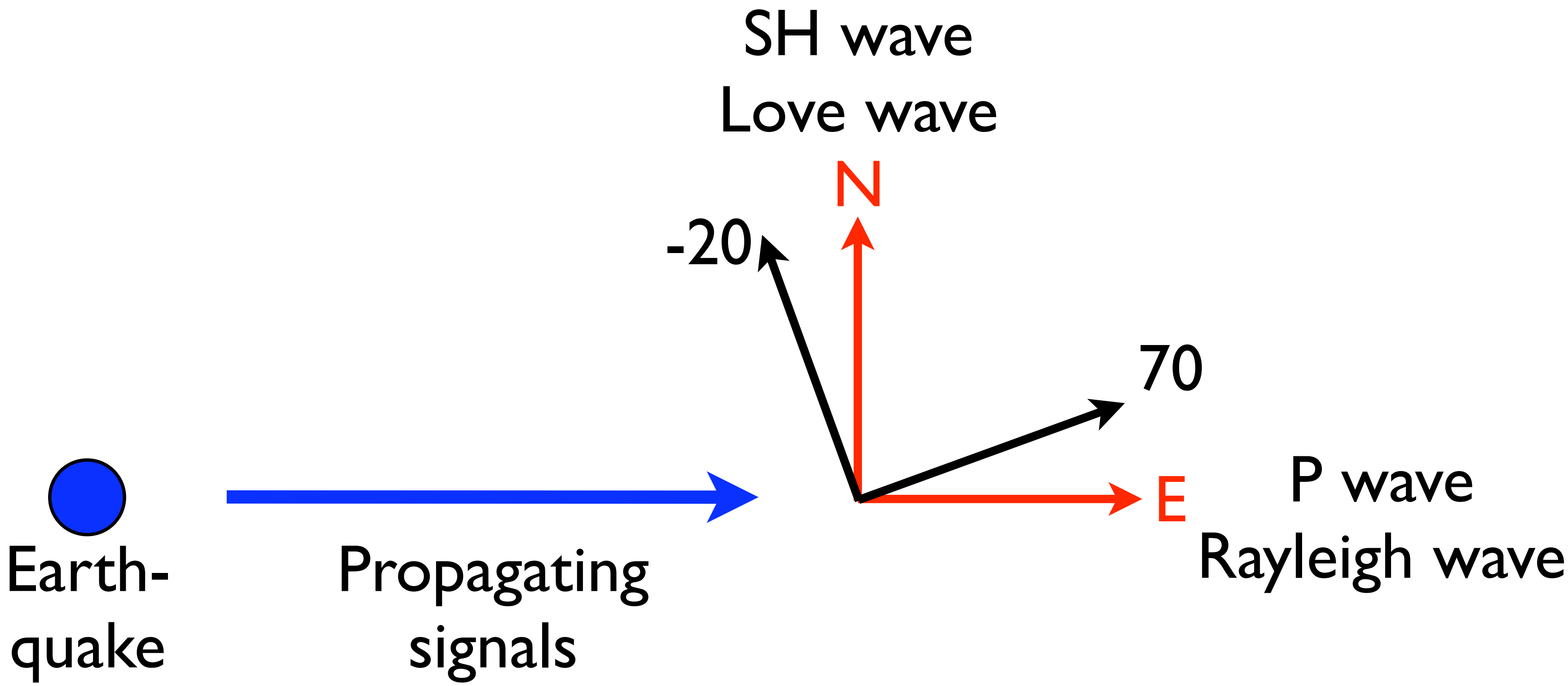
Horizontal Polarization Problems

Desired (assumed) orientation of seismometer

True orientation of seismometer



Natural Polarization of Earthquake Signals



SH wave
Love wave

N

-20

70

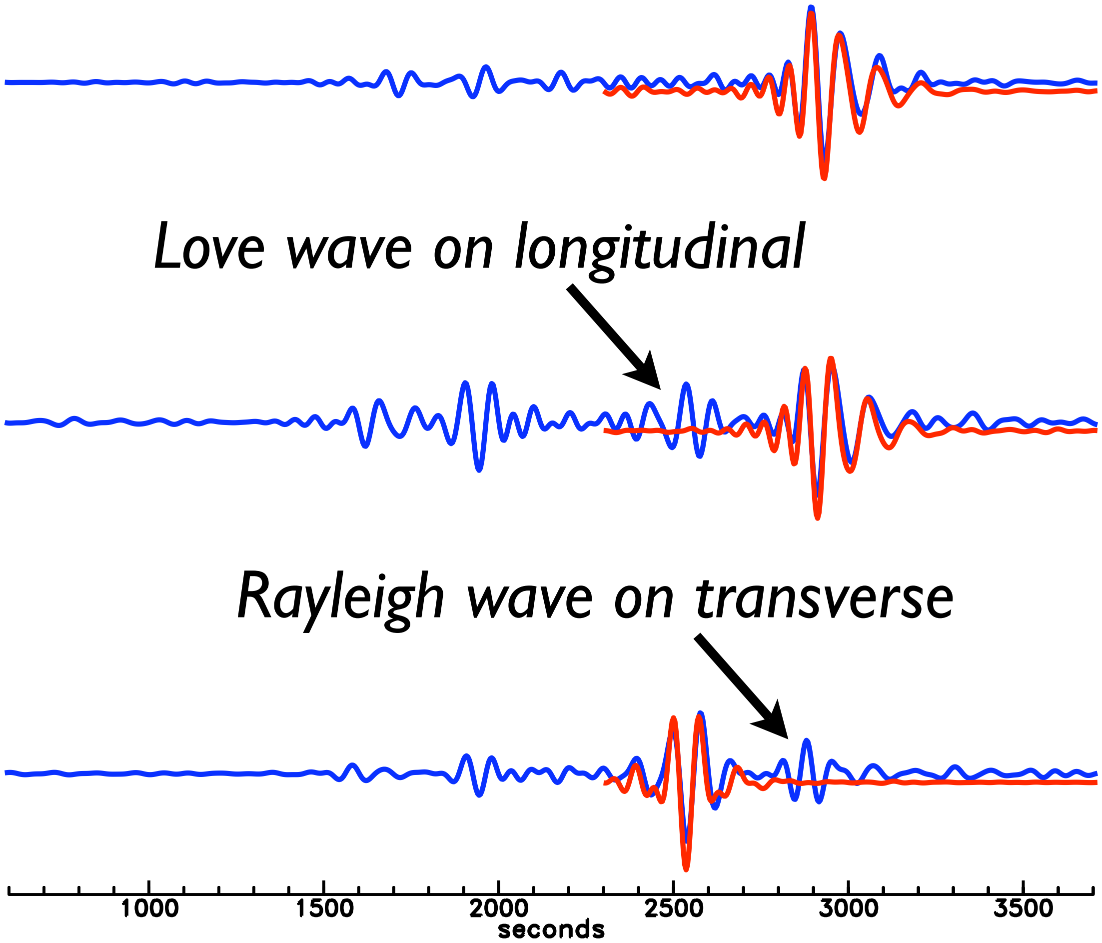
E

P wave
Rayleigh wave

Earth-
quake

Propagating
signals

Symptoms of a misoriented sensor



Vertical

Love wave on longitudinal

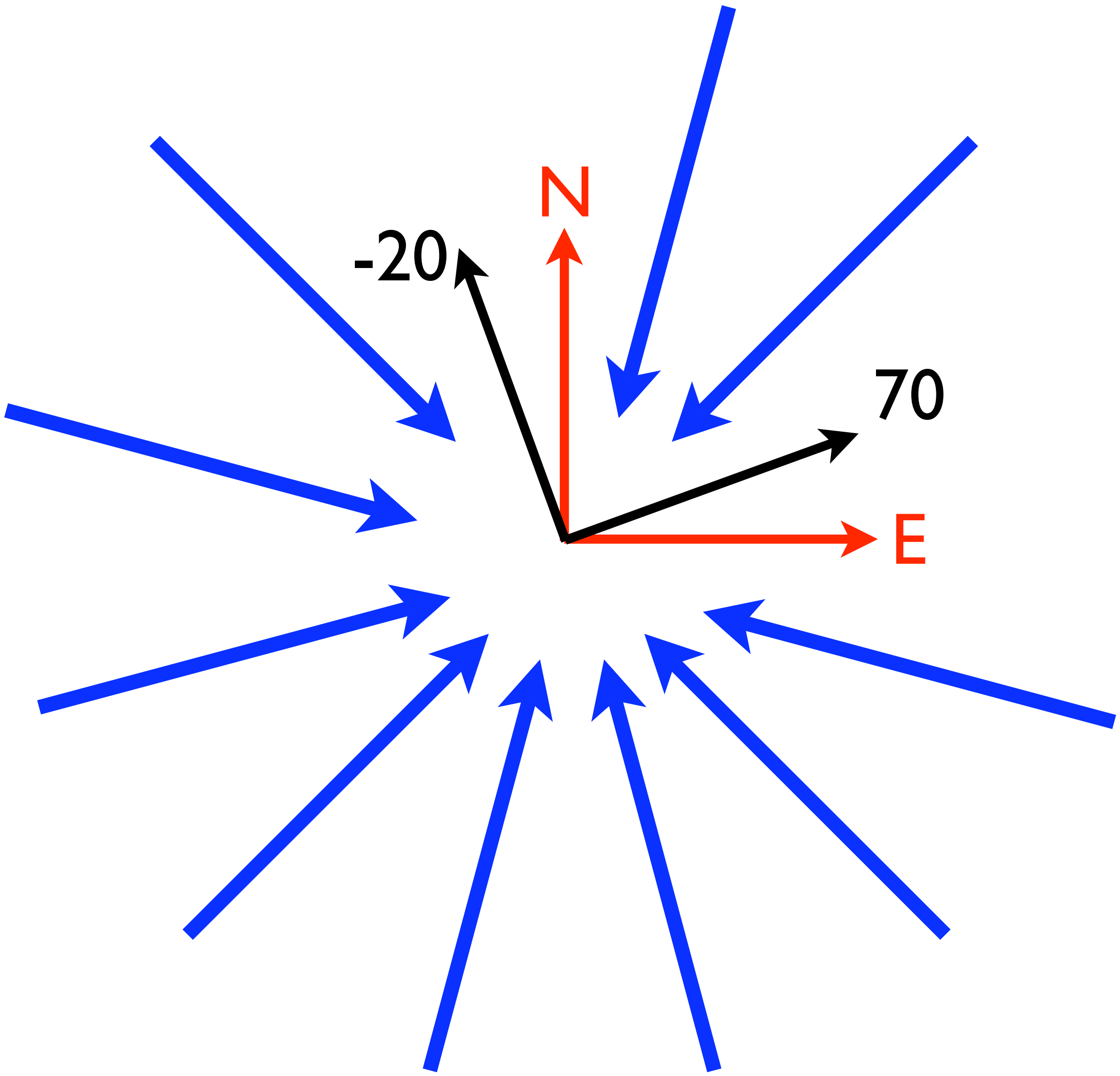
Longitudinal

Rayleigh wave on transverse

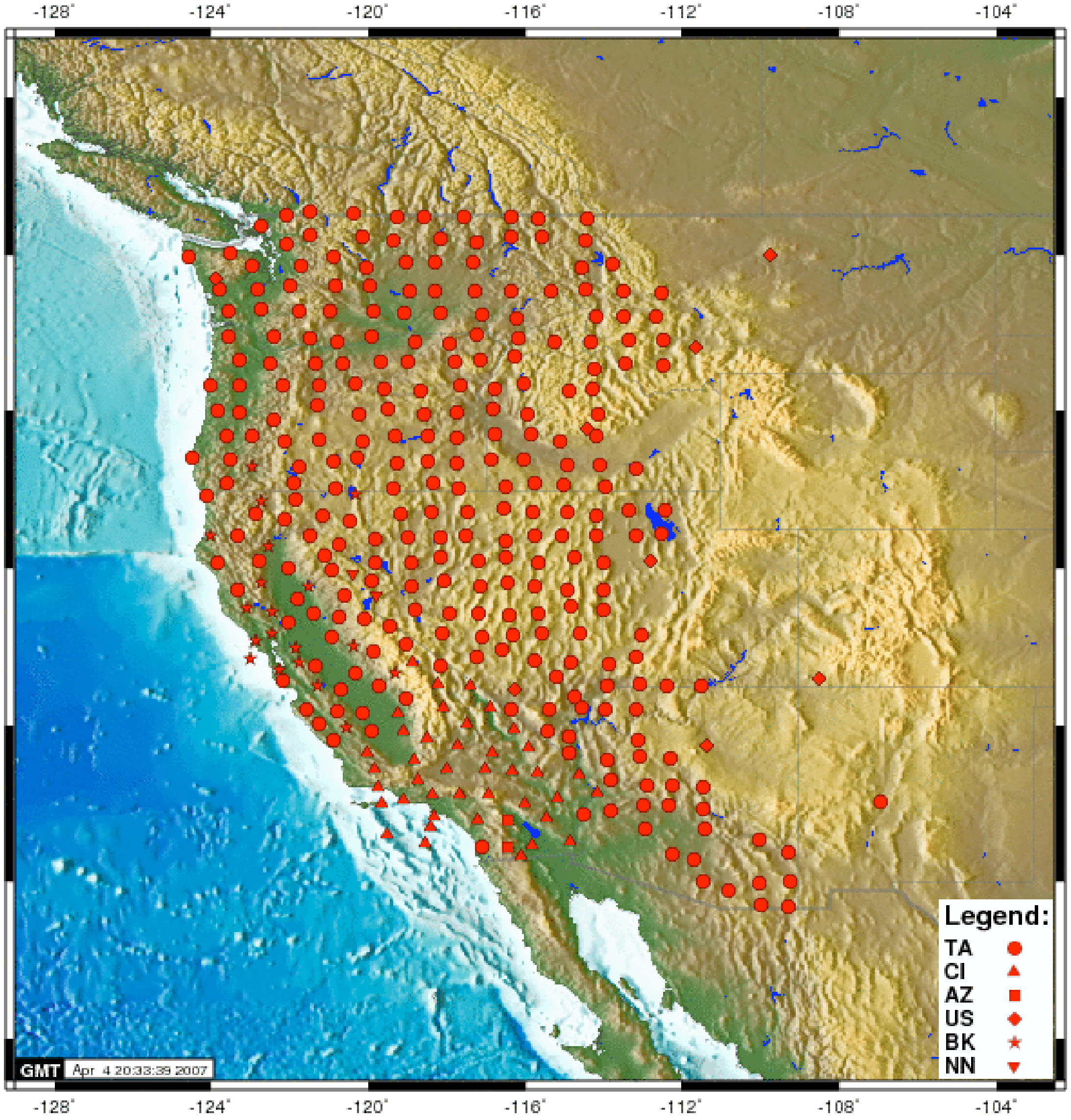
Transverse

Station D09A, earthquake on 08/20/2007

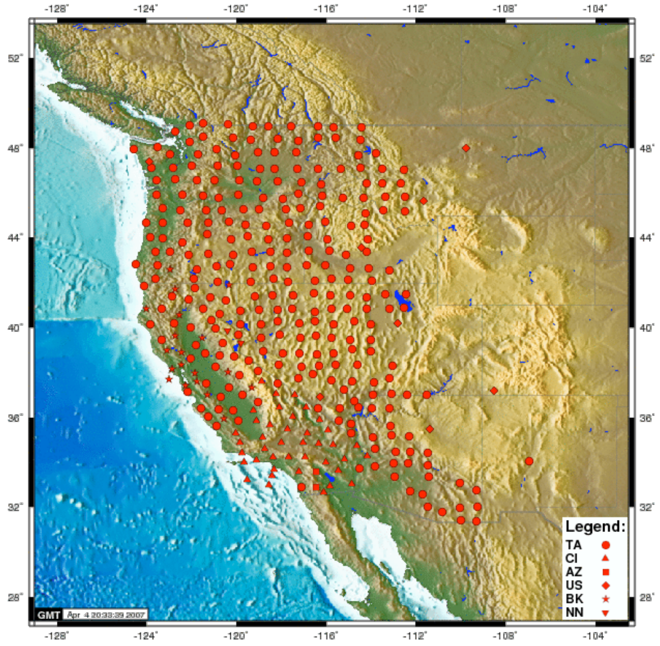
Many earthquake signals -- invert for orientation of sensor



USArray Transportable Array, April 2007



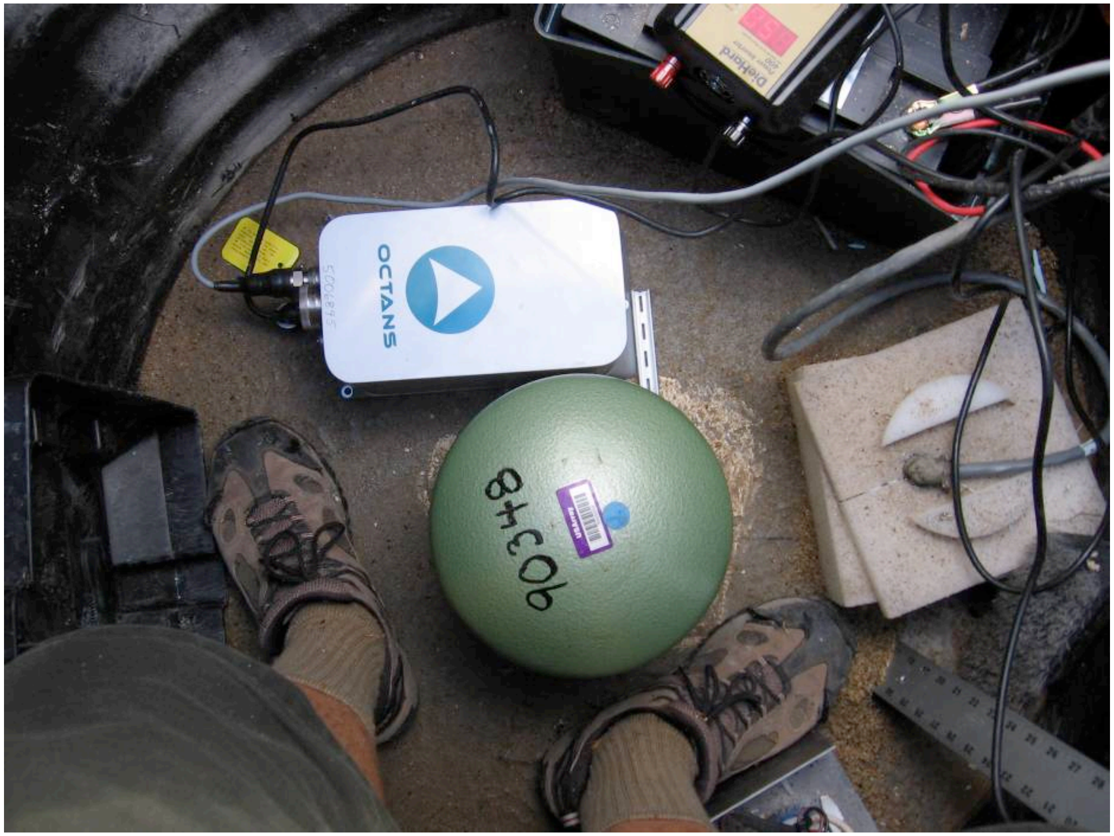
Polarization analysis of USArray data using earthquake signals recorded in 2006-2007



400+ USArray stations

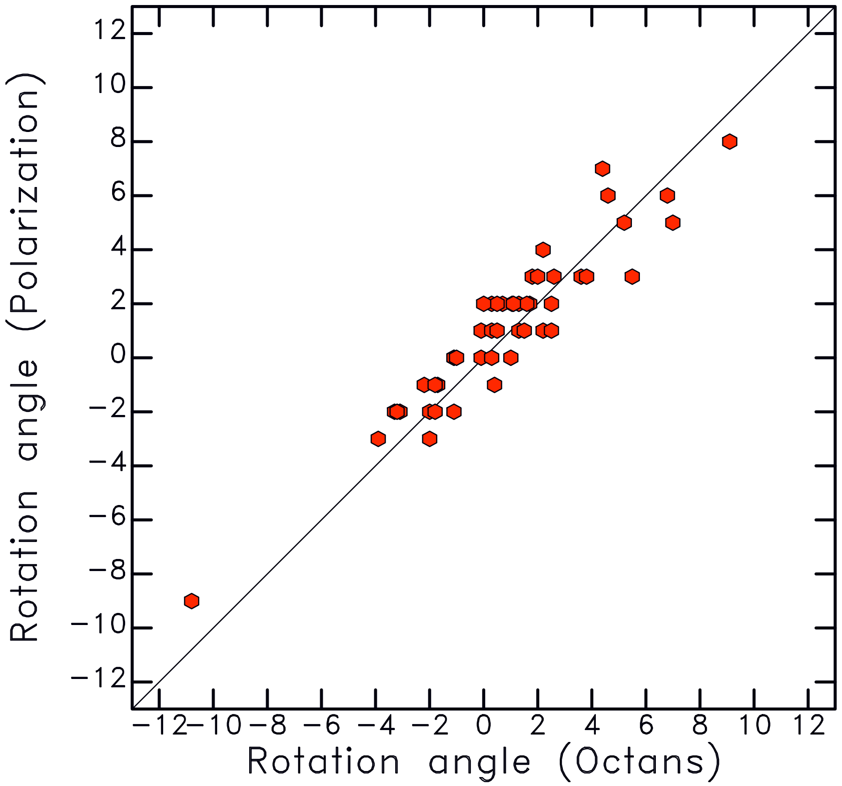
Result:
> 5% misoriented > 10 degrees
> 10 % misoriented > 5 degrees

Octans interferometric laser gyro



Agreement of field (Octans) and polarization angles

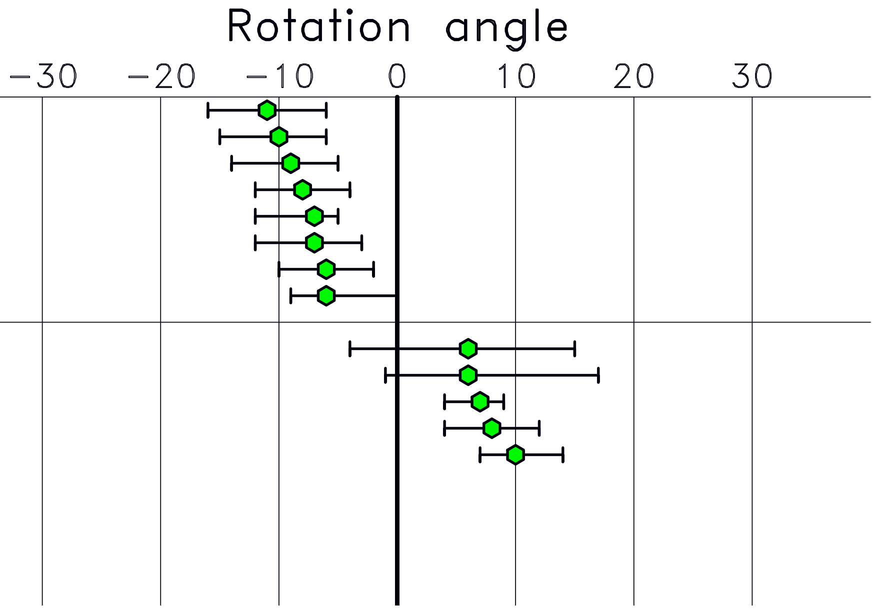
estimated from seismograms



measured in the field

Outliers (>5 deg)
II, IU, IC
as of 2009/11/08

KDAK-II	LHE-00	LHN-00	84.1	354.1	389	-11
MIDW-IU	LHE-00	LHN-00	90.0	0.0	55	-10
KIV-II	LHE-00	LHN-00	90.0	0.0	169	-9
AAK-II	LHE-00	LHN-00	90.0	0.0	232	-8
KMBO-IU	LHE-00	LHN-00	90.0	0.0	27	-7
TLY-II	LHE-00	LHN-00	90.0	0.0	301	-7
CASY-IU	LHE-00	LHN-00	270.0	180.0	45	-6
SNZO-IU	LH1-00	LH2-00	271.0	1.0	395	-6
LSA-IC	LHE-00	LHN-00	90.0	0.0	205	6
TIXI-IU	LHE-00	LHN-00	90.0	0.0	48	6
HRV-IU	LHE	LHN	90.0	0.0	75	7
PALK-II	LH1-00	LH2-00	6.5	96.5	202	8
SSPA-IU	LHE-00	LHN-00	90.0	0.0	198	10



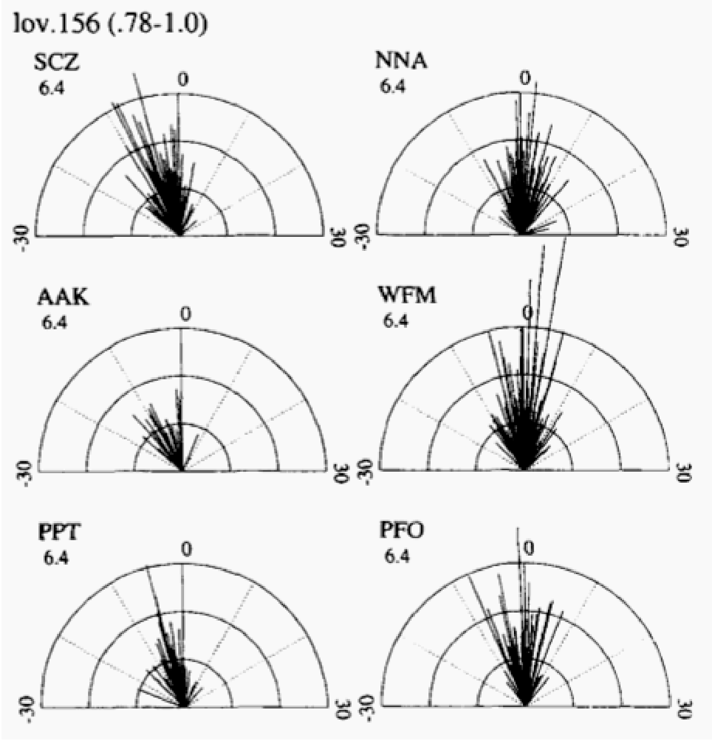
several GSN outliers have been eliminated in the last year or so by updates to metadata or (for secondary sensors) re-orientation of the sensor

Sensor orientation

Most GSN and USArray TA stations are well oriented, but not all.

Why does it matter?

- Modeling of earthquake sources
- Measurement of Love wave / toroidal mode parameters
- Estimates of anisotropy
- Estimates of off-great-circle arrival angle, for both elastic and anelastic structure

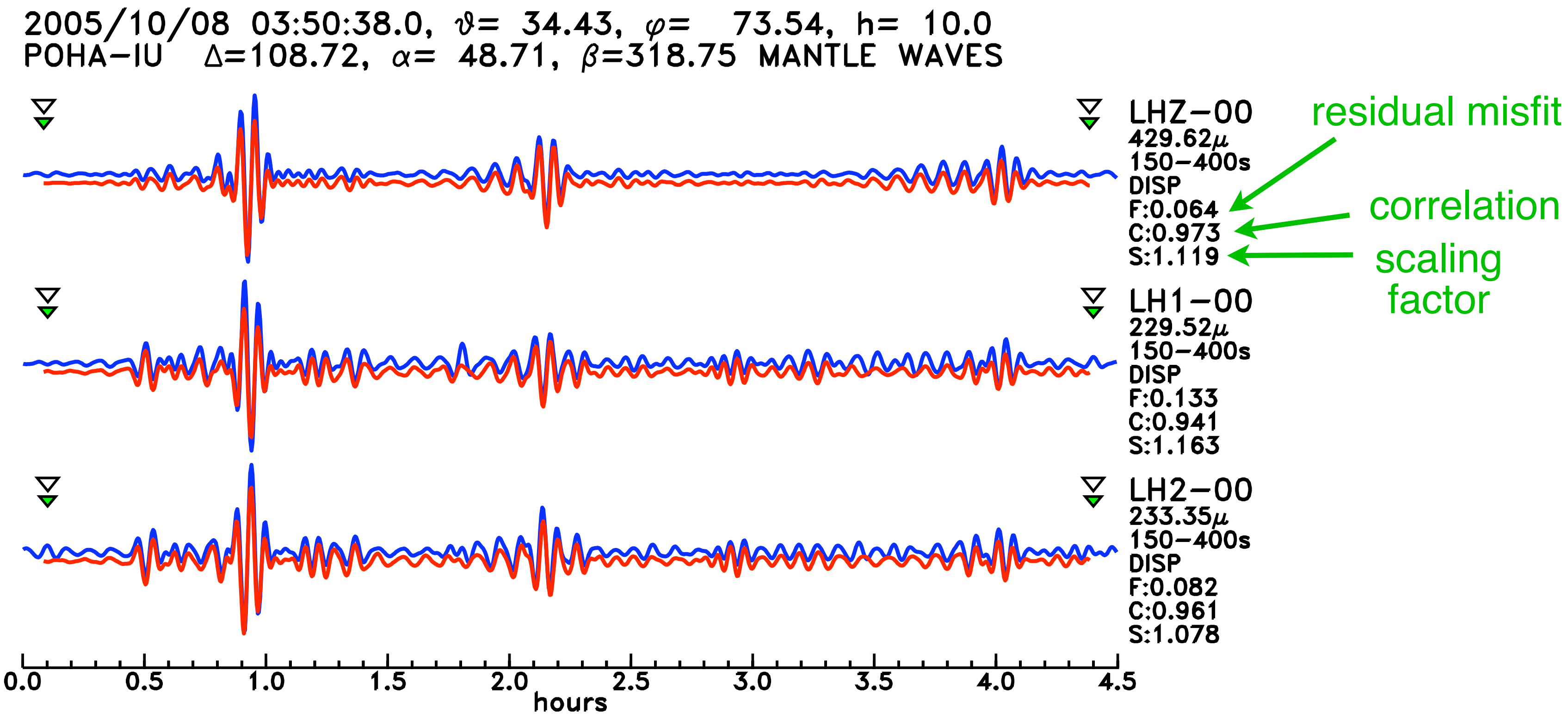


(Laske, 1995)

4b. Sensor response stability

Blue - observed seismograms

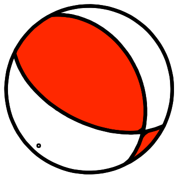
Red - synthetic seismograms



residual misfit

correlation

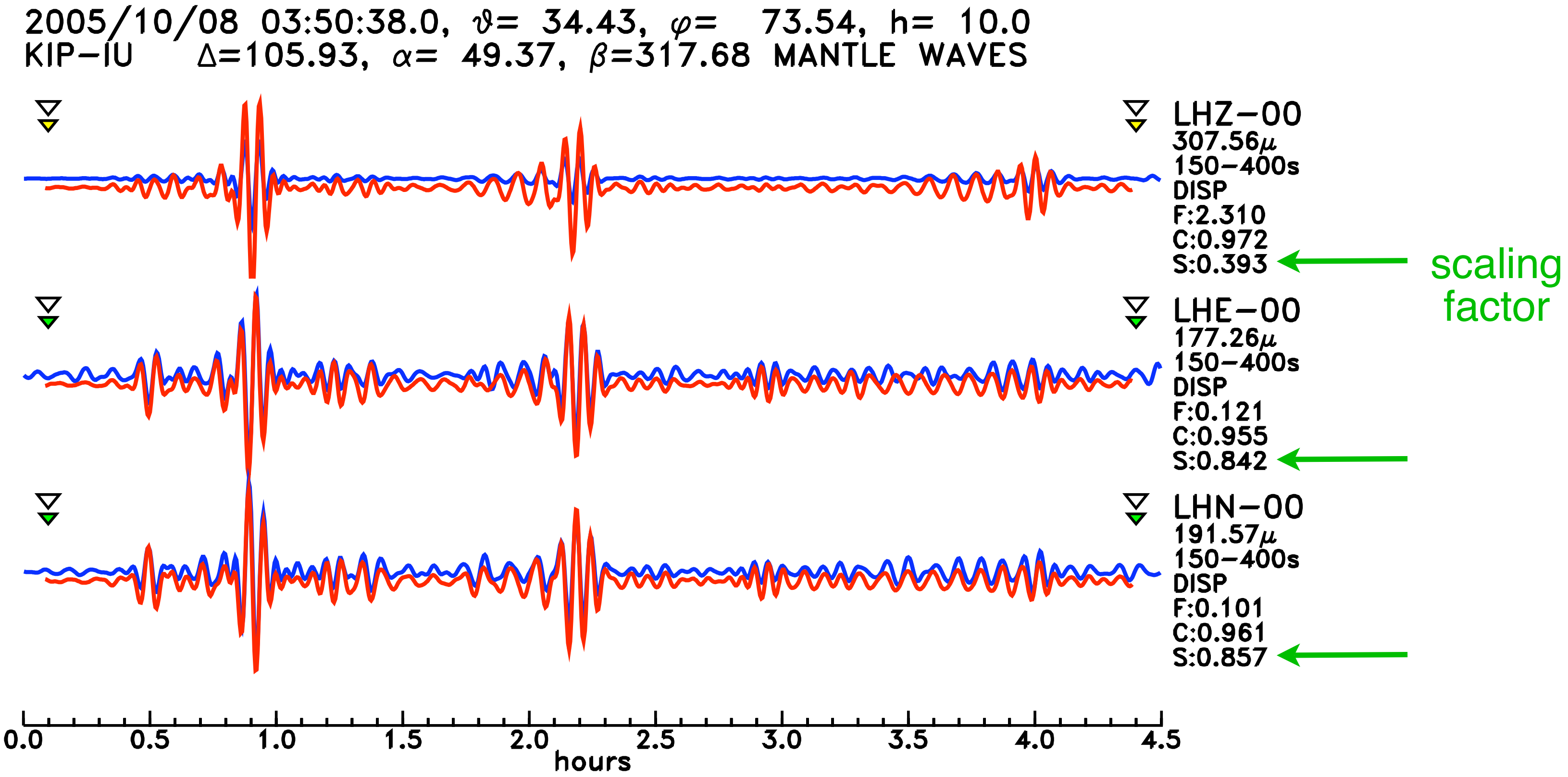
scaling factor



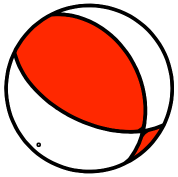
$$S = \frac{\sum_{i=1}^{N} o_i s_i}{\sum_{i=1}^{N} s_i^2}$$

Blue - observed seismograms

Red - synthetic seismograms

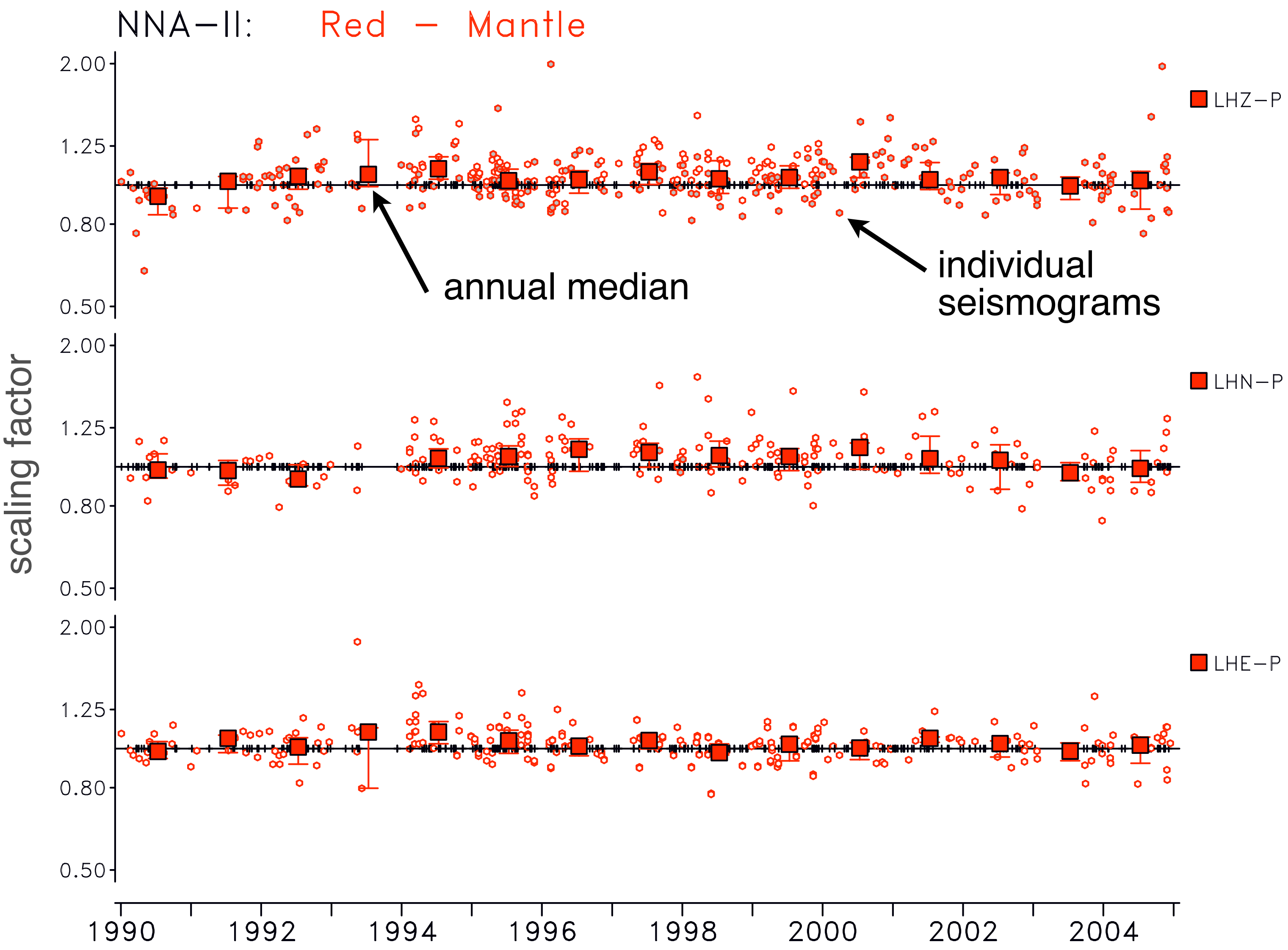


scaling factor



$$\boxed{S = \frac{\sum_{i=1}^{N} o_i s_i}{\sum_{i=1}^{N} s_i^2}}$$

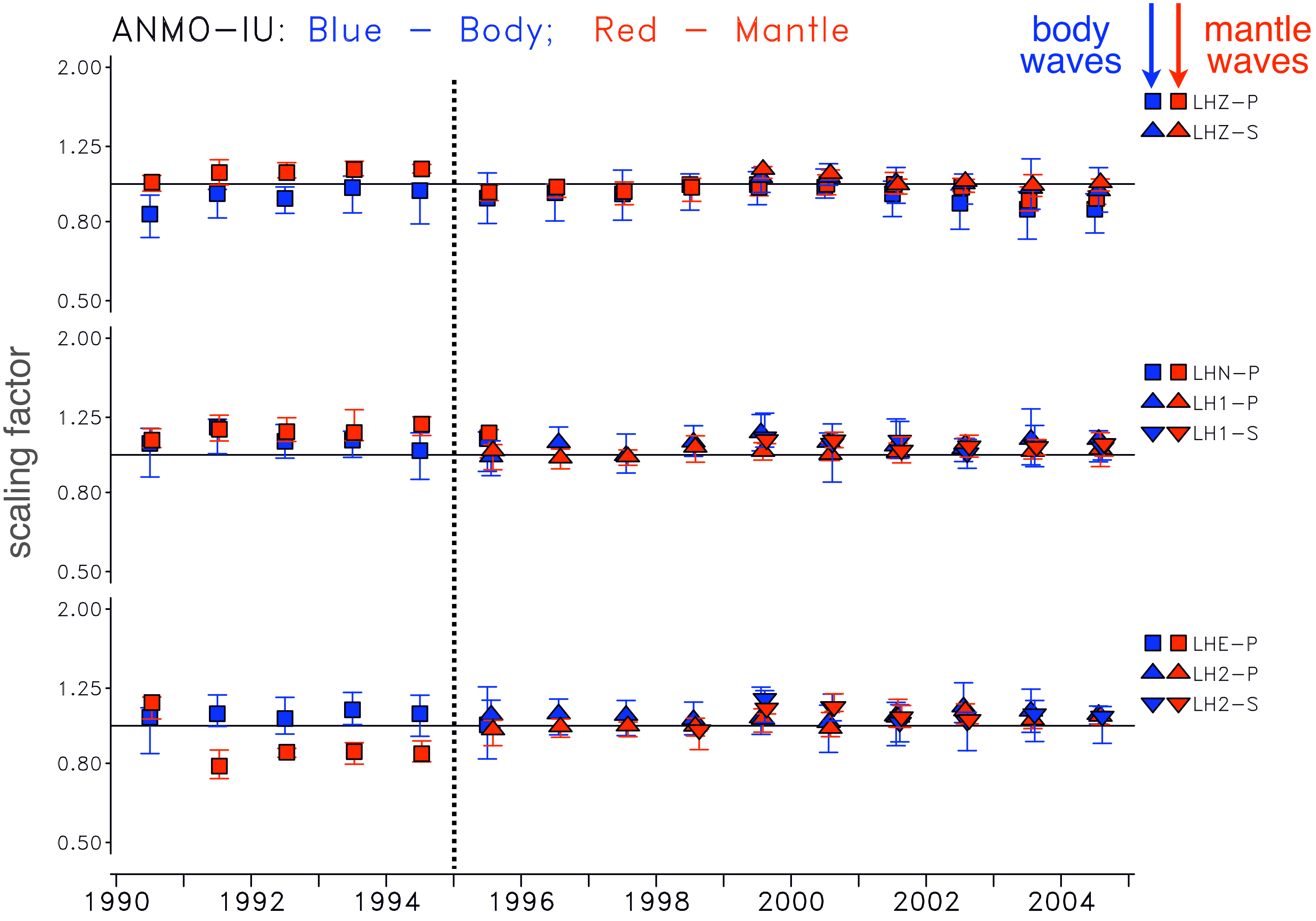
Scaling factors at NNA-II, 1990-2004



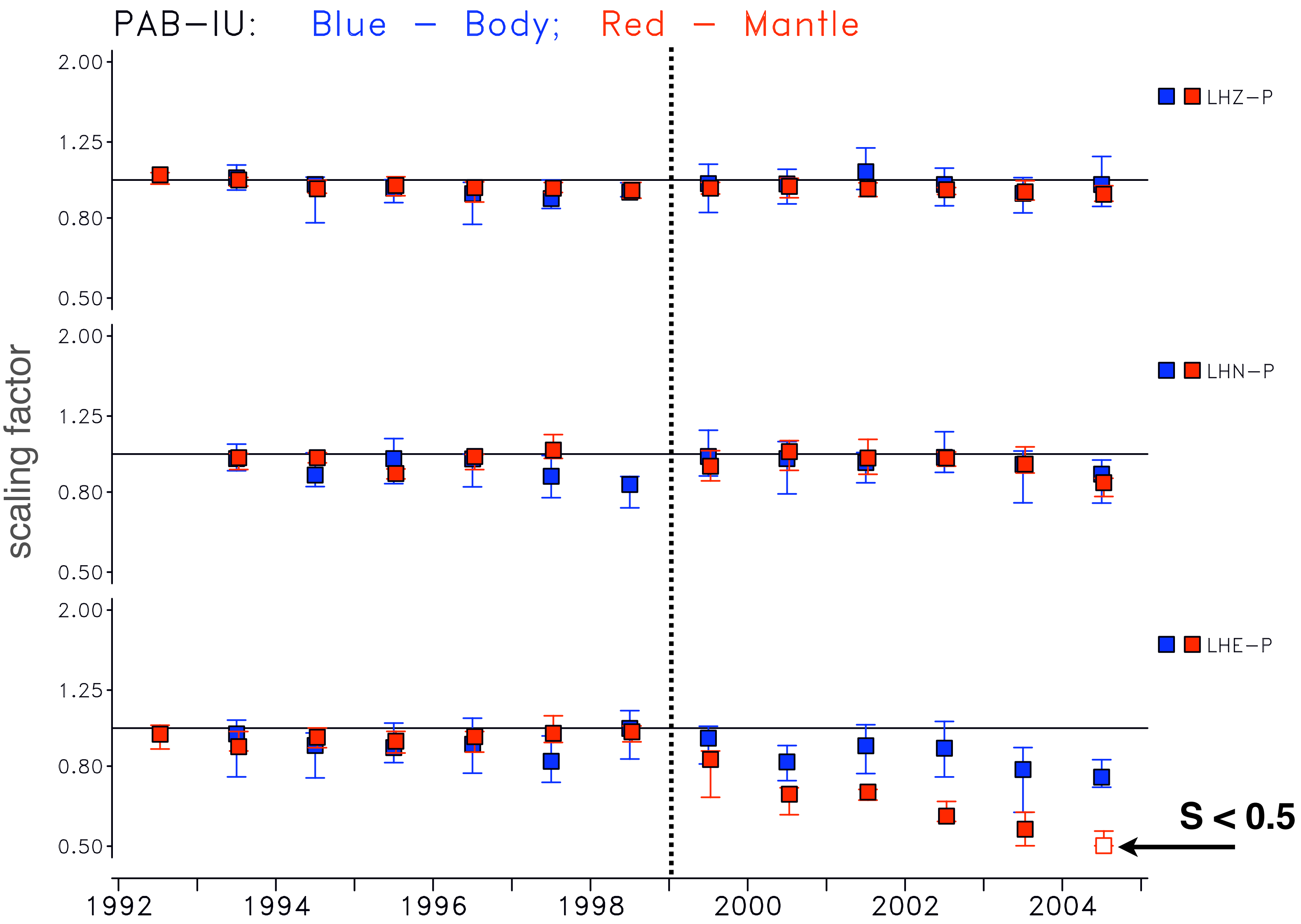
annual median

individual seismograms

Scaling factors at ANMO-IU, 1990-2004

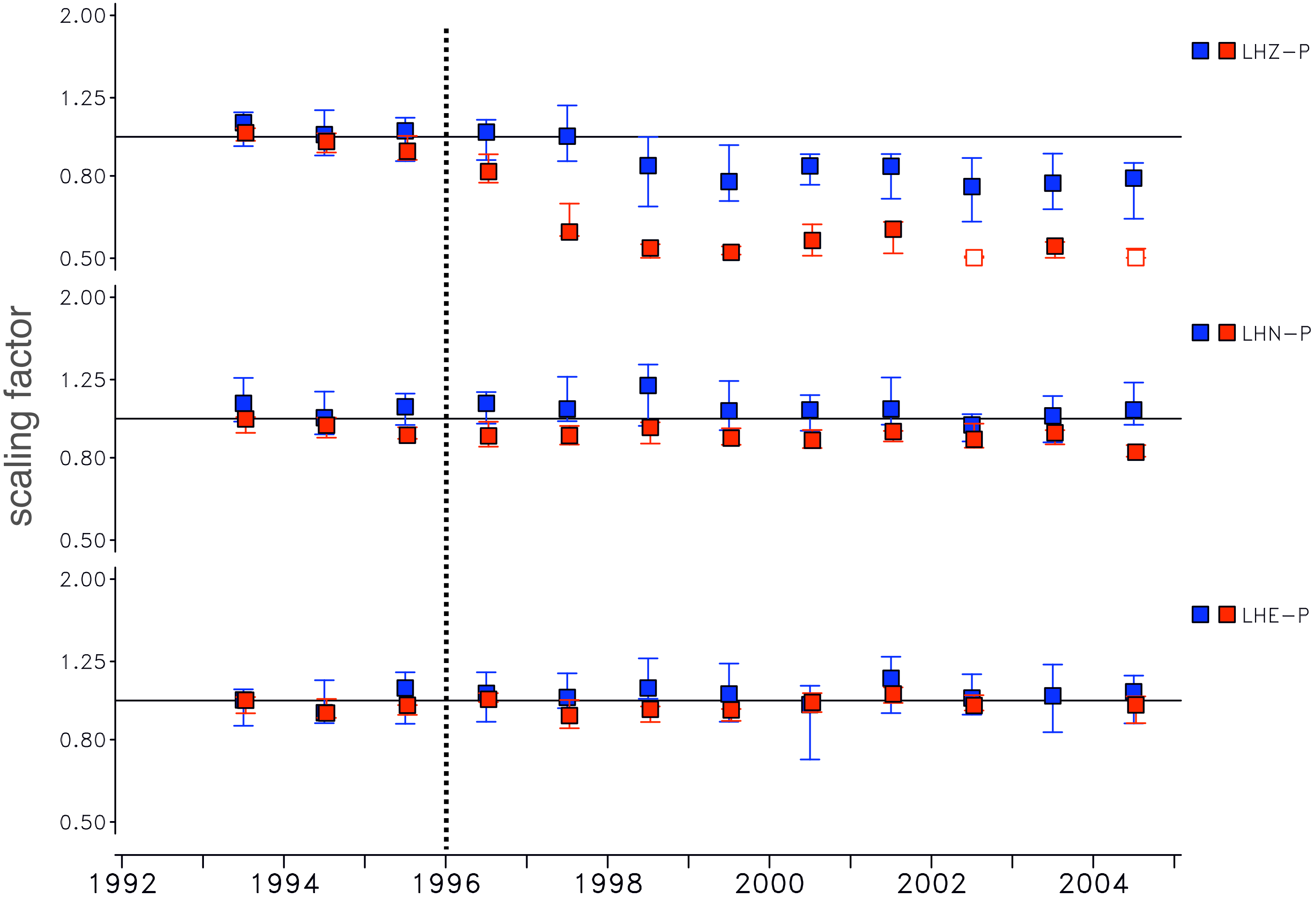


Scaling factors at PAB-IU, 1992-2004

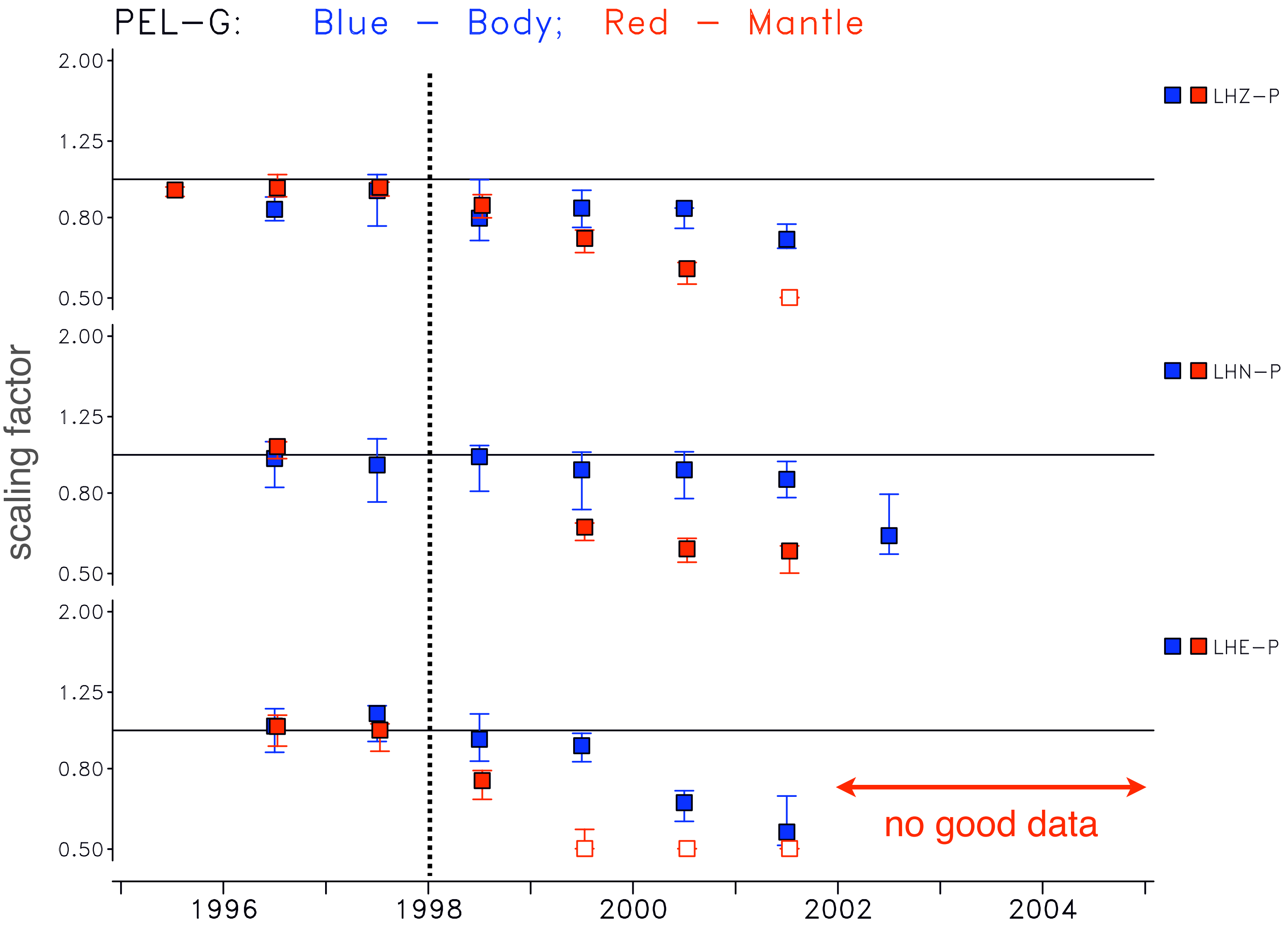


Scaling factors at LVZ-II, 1993-2004

LVZ−II: Blue − Body; Red − Mantle

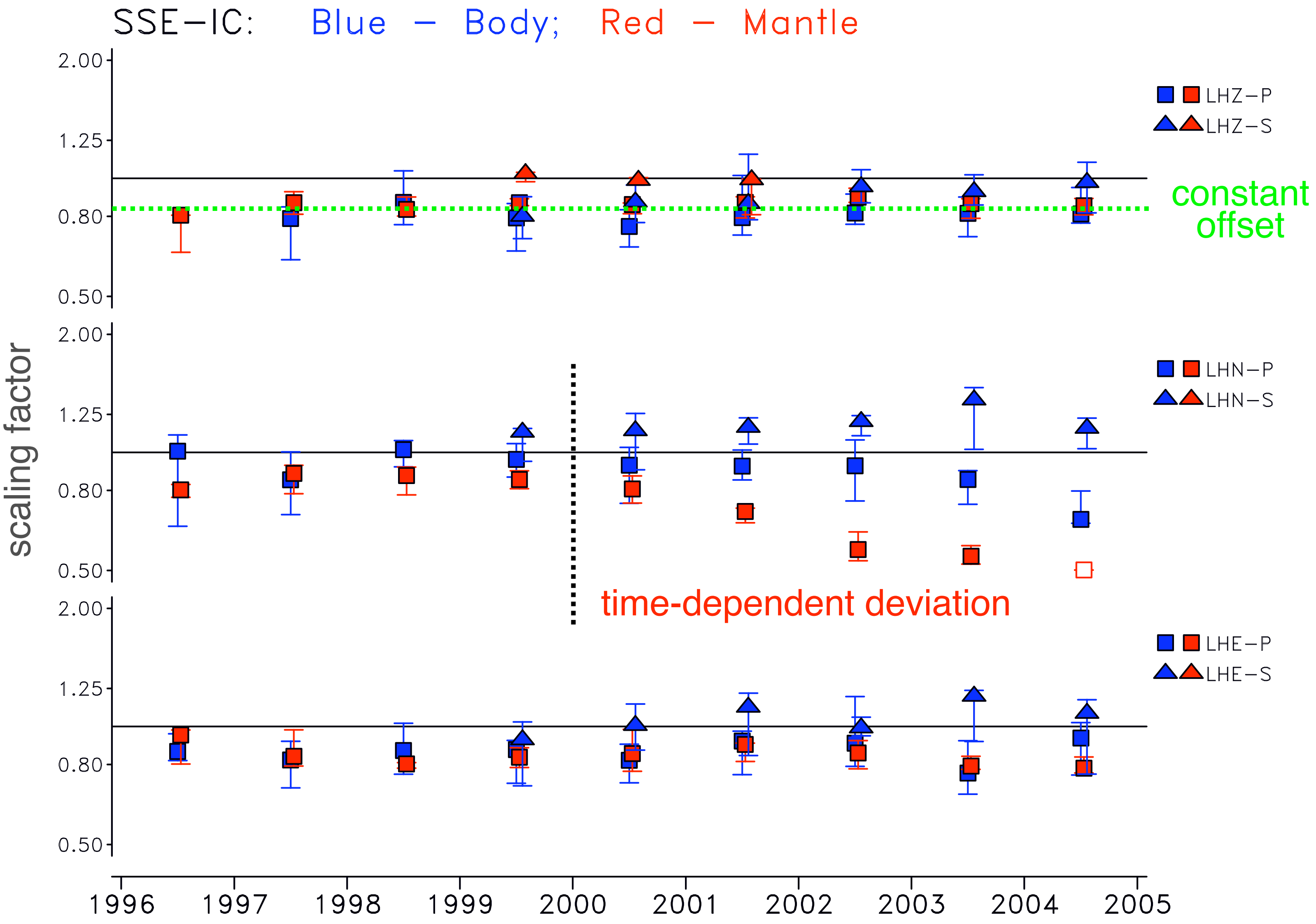


Scaling factors at PEL-G, 1996-2002



no good data

Scaling factors at SSE-IC, 1996-2004



Summary

- All results are available at: www.ldeo.columbia.edu/~ekstrom/Projects/WQC/SCALING
- Most stations show no, or small, deviations from the reported response
- A few stations (e.g., GTSN) show constant offsets in gain of 10-20%
- *Approximately 15% of stations equipped with STS-1 seismometers show a time- and frequency-dependent deterioration of the true gain*

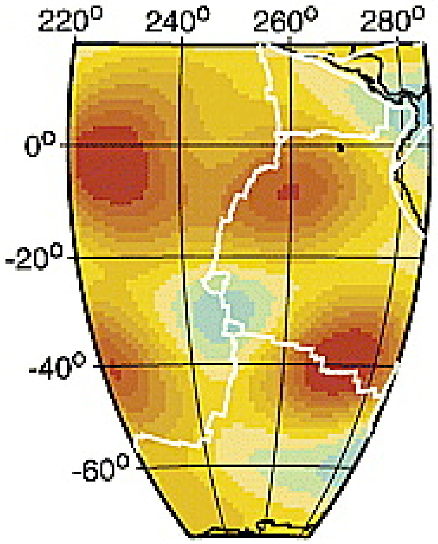
➡ Cause of problem?

- How to fix instruments?
- How to fix response information retroactively?

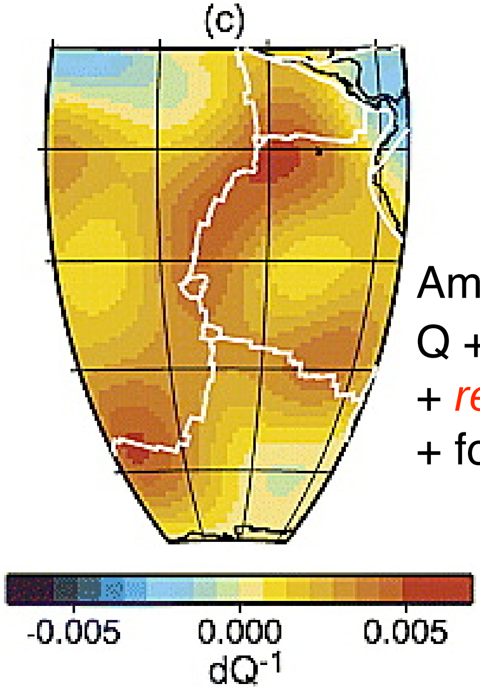
➡ Recommend regular instrument calibration

Why does it matter?

- Amplitudes carry critical information for improving models of elastic and inelastic (Q) structure
- Also important for improvements in source modeling



Amplitude → Q



Amplitude →
Q + source factor +
+ *receiver factor*
+ focusing

(Dalton and Ekström, 2006)

A simpler way to do this - if you have two instruments (A and B) in the same location:

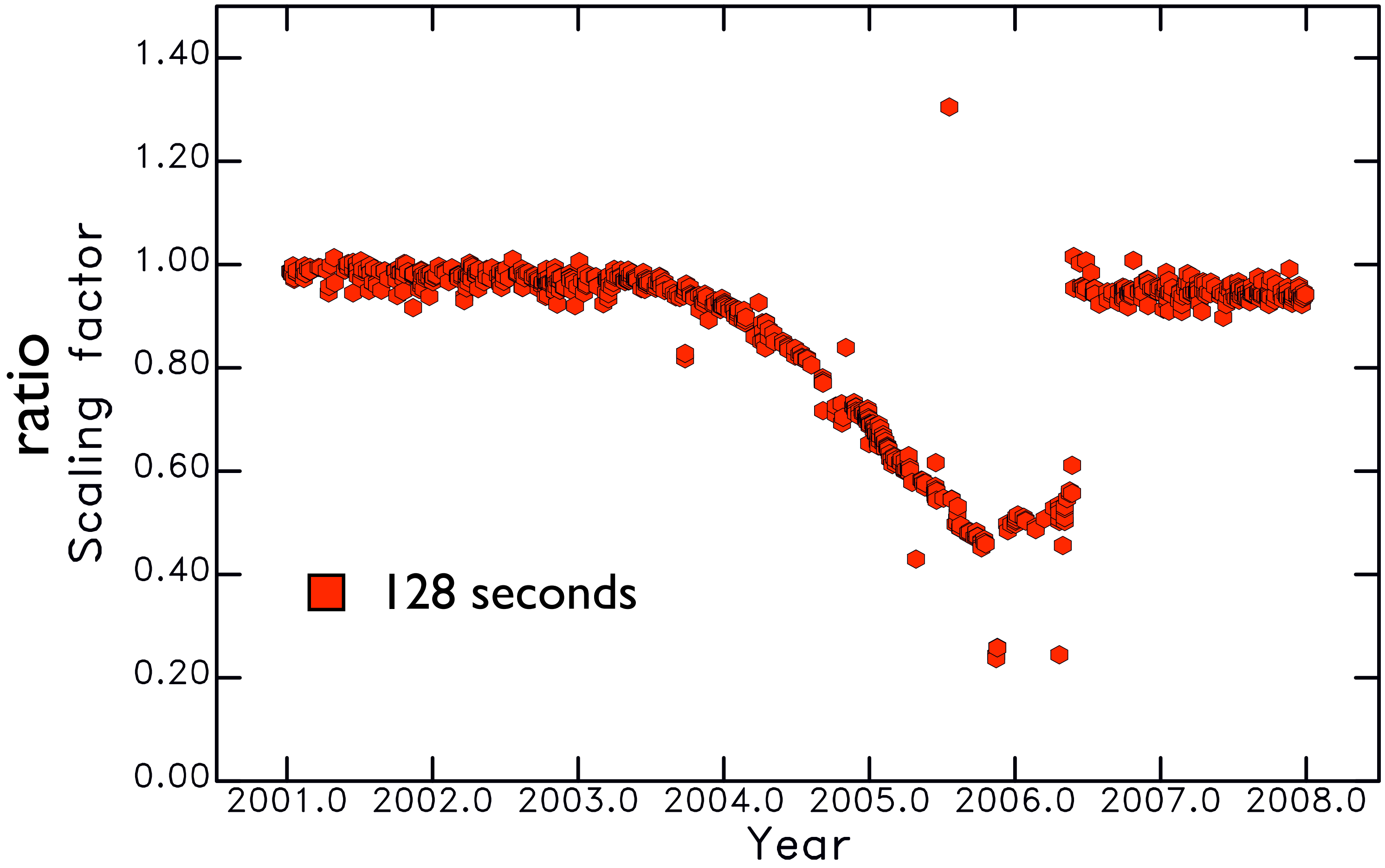
calculate ratio of displacements at some period during times of high signal coherence

$$\frac{\text{signal A}}{\text{response A}} = \text{displacement A} \qquad \text{(deconvolution)}$$

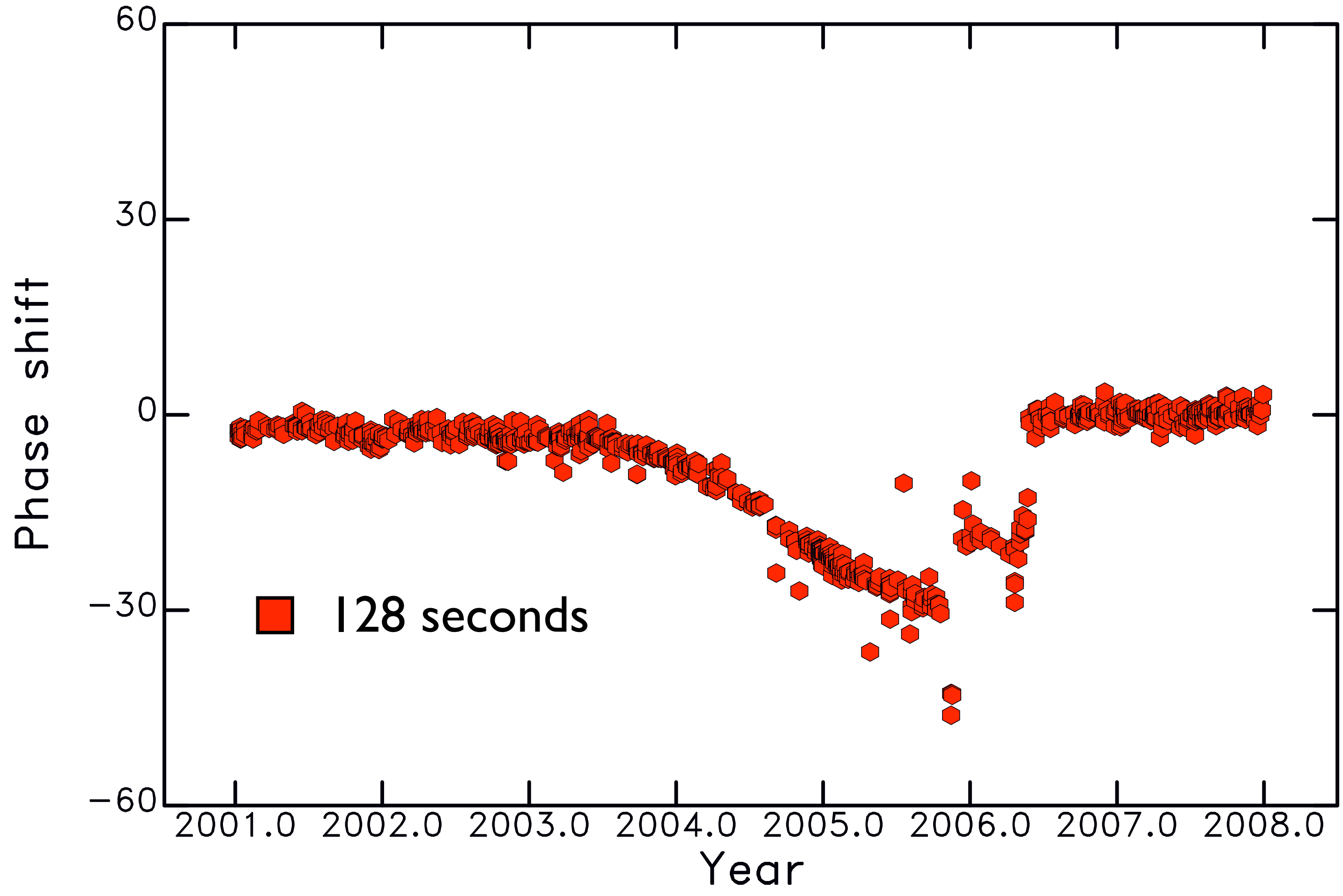
$$\frac{\text{signal B}}{\text{response B}} = \text{displacement B} \qquad \text{(deconvolution)}$$

$$\boxed{\text{ratio} = \frac{\text{displacement A}}{\text{displacement B}} \qquad \text{should be 1.0000!}}$$

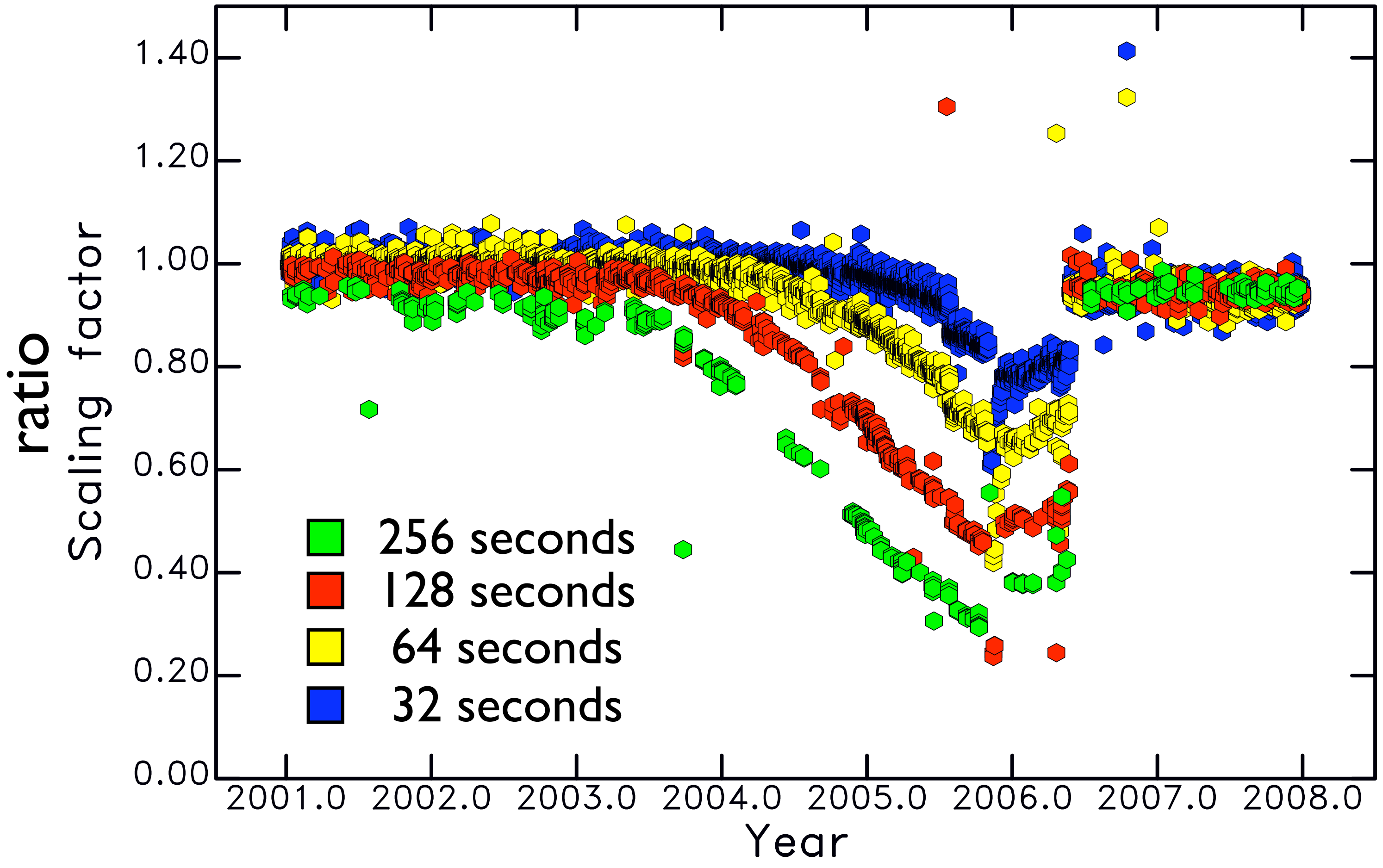
Scaling of STS-1 to STS-2 displacement at station KIP



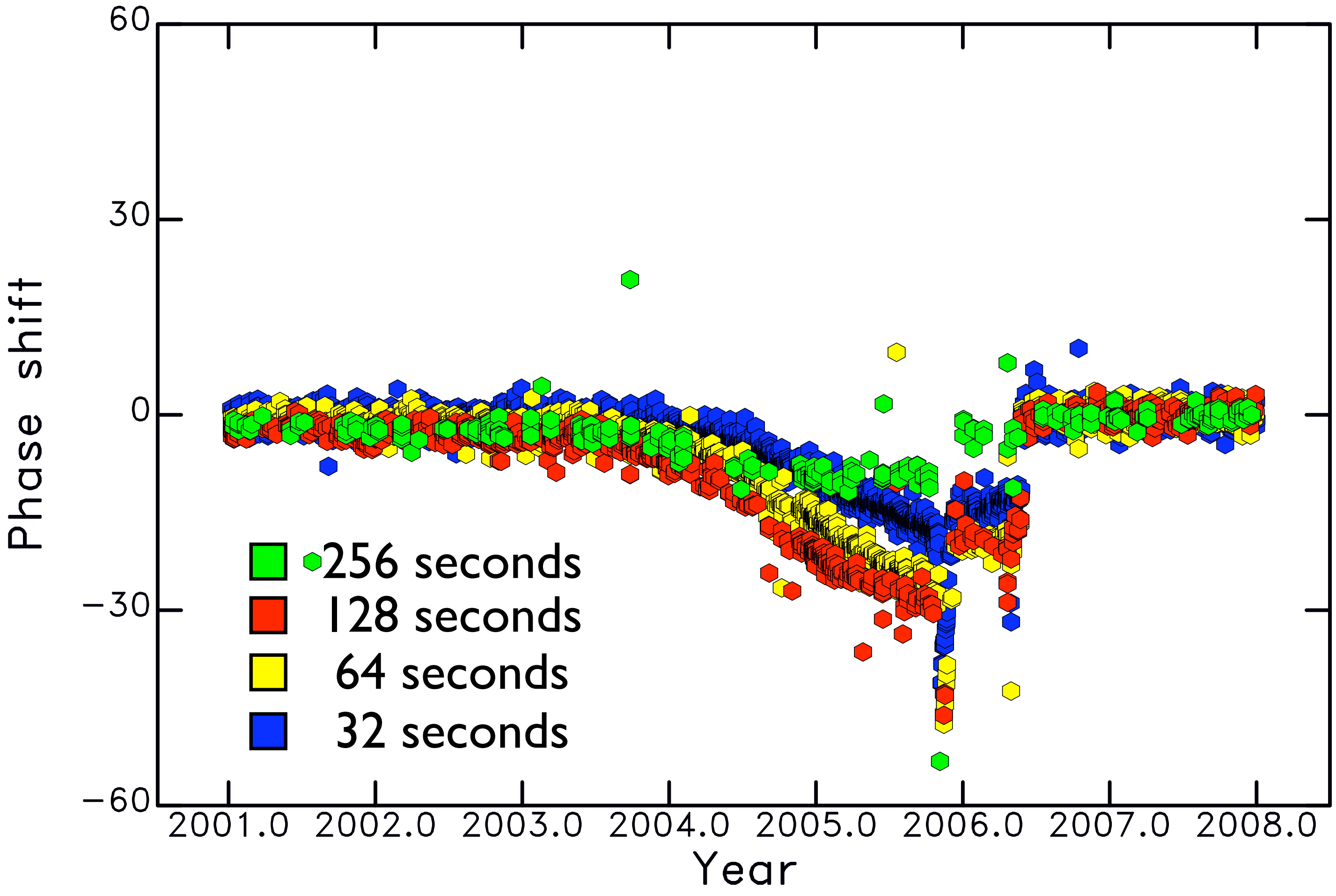
Phase shift of STS-1 to STS-2 displacement at station KIP



Scaling of STS-1 to STS-2 displacement at station KIP



Phase shift of STS-1 to STS-2 displacement at station KIP



Main points

1. The data can tell you a lot about your stations
2. Things change (calibrate!)
3. All networks can be improved

timing
orientation
response
noise level

All are important!